

# Collisional processes in extrasolar planetesimal discs – dust clumps in Fomalhaut’s debris disc

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## ABSTRACT

This paper presents a model for the outcome of collisions between planetesimals in a debris disc, and assesses the impact of collisional processes on the structure and size distribution of the disc. The model is presented by its application to Fomalhaut’s collisionally replenished dust disc; a recent 450- $\mu\text{m}$  image of this disc shows a clump embedded within it with a flux  $\sim 5$  per cent of the total. The following conclusions are drawn. (i) Spectral energy distribution modelling is consistent with Fomalhaut’s disc having a collisional cascade size distribution extending from bodies 0.2 m in diameter (the largest that contribute to the 850- $\mu\text{m}$  flux) down to 7- $\mu\text{m}$ -sized dust (smaller grains are blown out of the system by radiation pressure). (ii) Collisional lifetime arguments imply that the collisional cascade starts with planetesimals 1.5–4 km in diameter, and so has a mass of 20–30  $M_{\oplus}$ . Any larger bodies must be predominantly primordial. (iii) Constraints on the time-scale for the ignition of the collisional cascade from planet formation models are consistent with these primordial planetesimals having the same distribution as the cascade extending up to 1000 km, resulting in a disc mass of 5–10 times the minimum solar nebula mass. (iv) The debris disc is expected to be intrinsically clumpy, as planetesimal collisions result in dust clumps that can last up to 700 orbital periods. The intrinsic clumpiness of Fomalhaut’s disc is below current detection limits, but it could be detectable by future observatories such as *ALMA*, and could provide the only way of determining this primordial planetesimal population. Also, we note that such intrinsic clumpiness in an exozodiacal cloud-like disc could present a confusion limit when trying to detect terrestrial planets. (v) The observed clump could have originated in a collision between two runaway planetesimals, both larger than 1400 km in diameter. It appears unlikely that we should witness such an event unless both the formation of these runaways and the ignition of the collisional cascade occurred relatively recently (within the last  $\sim 10$  Myr), however this is a topic which would benefit from further exploration using planet formation and collisional models. (vi) Another explanation for Fomalhaut’s clump is that  $\sim 5$  per cent of the planetesimals in the ring were trapped in 1:2 resonance with a planet orbiting at 80 au when it migrated out as a result of the clearing of a residual planetesimal disc. The motion on the sky of such a clump would be 0.2 arcsec  $\text{yr}^{-1}$ , and it would be more prominent at shorter wavelengths.

**Key words:** circumstellar matter – stars: individual: Fomalhaut – planetary systems: formation.

## 1 INTRODUCTION

At least 15 per cent of main-sequence stars exhibit detectable far-infrared (far-IR) emission in excess of that expected from the photosphere (Plets & Vynckier 1999; Lagrange, Backman & Artymowicz Lagrange et al. 2001). The spectral energy distribution (SED) of this

excess implies that it is thermal emission from dust in regions analogous to the Kuiper belt (i.e. at  $>30$  au from the star). The few stars for which this emission has been imaged confirm this location for the dust and show that it is confined to narrow ring-like discs (e.g. Holland et al. 1998; Greaves et al. 1998; Jayawardhana et al. 1998). The short lifetime inferred for this dust, due to both mutual collisions and radiation forces, implies that it must be continually replenished, and it is thought that the dust disc is fed by the collisional grinding-down of a population of larger planetesimals which

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may have formed during the planetary formation phase of the system (Backman & Paresce 1993). Furthermore, it has often been speculated that the absence of dust close to the star in these systems has been caused by the clearing of the residual planetesimals by a planetary system.

The debris disc images also show that the dust rings are neither smooth nor symmetrical: recent submillimetre (submm) images of the Fomalhaut disc show a clump containing 5 per cent of the total disc flux (Holland et al. 2002); there are several clumps in the  $\epsilon$  Eridani disc, the largest of which contains 6–7 per cent of the disc flux (Greaves et al. 1998); a 5 per cent brightness asymmetry is seen in the structure of the HR 4796 disc (Telesco et al. 2000), which could be accounted for by a clump containing  $\sim 0.5$  per cent of the disc flux; Vega’s dust disc is dominated by emission from two dust clumps (Holland et al. 1998; Koerner, Sargent & Ostroff 2001); and  $\beta$  Pictoris exhibits several small clumps in the mid-plane of the north-east extension of its optical disc (Kalas et al. 2000), as well as a submm emission peak lying  $\sim 34$  arcsec south-west of the star (Holland et al. 1998; Zuckerman 2001).

While clumps are now known to be a common feature of debris discs, their origin, and whether this is the same in all cases, is still unclear. One way to form clumps is the gravitational perturbations of unseen planets, as we know that their effect on the orbits of planetesimals in these discs could cause the observed dust rings to contain both offset and warp asymmetries (Wyatt et al. 1999, hereafter WDT99) as well as resonant clumps (Wyatt 1999; Ozernoy et al. 2000). Indeed, the observed asymmetries have been modelled by various authors and shown to be possible evidence of planets of mass between  $10 M_{\oplus}$  and  $2 M_J$  orbiting close to the inner edge of the rings (typically  $\sim 30$  au). However, it is difficult to test this hypothesis as the proposed planets are not detectable with current techniques, and such interpretations must be treated with caution until alternative explanations have been explored.

Another mechanism for forming clumps in these discs is stochastic collisions between their largest planetesimals. A collisionally produced disc is inherently clumpy because collisional debris follows the orbit of the parent planetesimal until it has had a chance to precess around that orbit. The question is, do collisions of a magnitude large enough to produce an observable dust clump happen often enough, and do the resulting clumps last long enough, to make a disc appear clumpy? The possibility of observing collisional clumps was first discussed by Stern (1996), who modelled the expected clumpiness of our own Kuiper belt, but the application to extrasolar debris discs has received little attention, as it was thought that collisions large enough to produce observable dust clumps would occur too infrequently, even if large enough planetesimals do exist in the discs. However, recent modelling of the asteroid belt stressed the importance of stochastic collisions by showing that the collisional destruction of asteroids large enough to more than double the brightness of the asteroid belt should occur about once every 20 Myr (Durda & Dermot 1997; see Fig. 19 of Grogan, Dermott & Durda 2001). This possibility may be supported by evidence of stochastic increases in the interplanetary dust flux from  $^3\text{He}$  in deep-sea sediments (Farley 1995). Also, dust clumps caused by collisions in the asteroid belt may already have been observed – the temporary dust tail activity of the asteroid–comet Elst/Pizarro may have resulted from collisions with other main-belt asteroids (Toth 2000). This prompted Wyatt et al. (2000) to revisit the collisional hypothesis by applying a simple collisional model to the recently obtained image of the Fomalhaut disc (Holland et al. 2002). This implied that collisions were unlikely to be causing Fomalhaut’s clump, but they concluded that this possibility could

not be ruled out because of uncertainties in the size distribution of the disc.

It is the purpose of this paper to explore in depth the role of collisions in shaping the structure of the Fomalhaut disc. In particular, the aim is to determine whether the clump observed by Holland et al. (2002) could be collisional in origin. One reason why it is crucial to test the collisional hypothesis is that the only observation that has been proposed to test whether the dust clumps are associated with planetary resonances relies on the fact that such clumps would orbit the star with the planet – motion that should be detectable on time-scales of a few years (Ozernoy et al. 2000). However, it may not be possible to use detection of such motion as an unambiguous test for the presence of planets, as clumps that are collisional in origin would also orbit the star. After summarizing the observations of the Fomalhaut disc (Section 2), we consider the size distribution expected from theoretical arguments (Section 3), then estimate this distribution observationally by modelling the observed SED (Section 4). A model of collisions is then developed which is used to explore the role of collisional destruction in forming the observed size distribution (Section 5). This model is then expanded so as to determine what magnitude of collision we expect to see in Fomalhaut’s disc due to the break-up of its planetesimals (i.e. its intrinsic clumpiness; Section 6). Finally, we explore the possible scenarios in which collisions might have produced Fomalhaut’s dust clump and propose another possible origin for the clump (Section 7). The conclusions are given in Section 8.

## 2 THE FOMALHAUT DISC

Fomalhaut was one of the first four main-sequence stars found to exhibit infrared emission in excess of that from the photosphere and one of only a handful that have had that emission resolved. It is an A3V star at a distance of  $R_* = 7.7$  pc that is estimated to have an age of  $t_{\text{sys}} = 200 \pm 100$  Myr, based on both the evolutionary isochrone models (Lachaume et al. 1999; Song et al. 2001) and the ages of the Castor moving group with which it is associated (Barrado y Navascues 1998) and of its common proper motion companion GL879 (Barrado y Navascues et al. 1997). The stellar parameters that have been adopted in this work are:  $T_{\text{eff}} = 9060$  K,  $\log g = 4.29$ ,  $M = 2 M_{\odot}$  (Song et al. 2001) and  $L = 13 L_{\odot}$  (Backman & Paresce 1993). In all modelling, the stellar spectrum is assumed to be that of a Kurucz model atmosphere with solar metallicity and the above parameters. No gas disc has been observed around Fomalhaut (Liseau 1999).

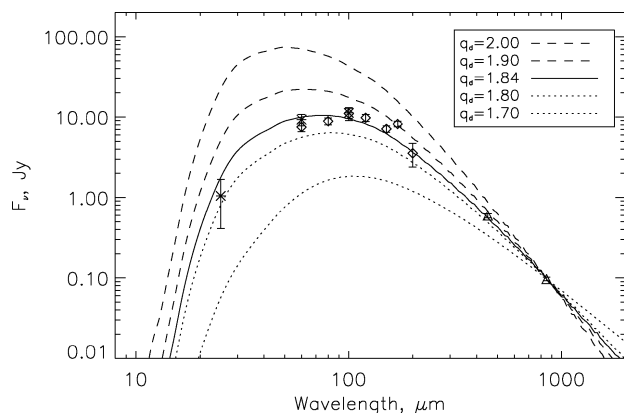
The photometric observations of emission from the Fomalhaut disc are summarized in Table 1, and this SED is plotted in Fig. 1. Recently the structure of the disc was mapped at both 450 and 850  $\mu\text{m}$  using SCUBA at the James Clerk Maxwell Telescope (JCMT) (Holland et al. 2002 Fig. 2a). The double-lobed emission feature seen in the images implies that the disc is being observed close to edge-on and that the dust is constrained to a narrow ring  $\sim 150$  au from the star. This is consistent with the SED of the disc, which implies emission from cool dust at a single temperature, as well as with previous observations of the structure of the disc (Holland et al. 1998; Harvey & Jefferys 2000). The images at both wavelengths also show an asymmetry in the brightness distribution. The spatial structure of the disc seen in the 450- $\mu\text{m}$  image (Fig. 2a) was modelled in Holland et al. (2002), where they showed that the observation is consistent with emission from a smooth axisymmetric ring (Fig. 2b), embedded within which is a bright clump with a flux equal to  $\sim 5$  per cent of the total disc flux ( $\sim 30$  mJy; Fig. 2c). Holland et al. (2002) argue that this clump is real and associated with the star, as

**Table 1.** Photometric observations of the excess emission from Fomalhaut,  $F_\nu$ , and the  $1\sigma$  errors,  $dF_\nu$ . The photospheric contribution has been subtracted from the observed emission using a Kurucz model atmosphere scaled to  $K = 0.992$  (Bouchet, Manfroid & Schmider 1991). The *IRAS* fluxes were taken from the Faint Source Catalog and a colour correction applied based on the temperature of the excess at each wavelength. This temperature is not well known at  $25\ \mu\text{m}$  and the resulting uncertainty in the derived flux is included in the errors. The *ISO* fluxes were taken from Walker et al. (in preparation).

$\lambda$ , $\mu\text{m}$	$F_\nu$ , Jy	$dF_\nu$ , Jy	Reference
25	1.24	0.6	<i>IRAS</i>
60	9.80	0.62	<i>IRAS</i>
60	7.51	0.92	<i>ISO</i>
80	8.86	0.87	<i>ISO</i>
100	10.95	0.66	<i>IRAS</i>
100	11.04	2.0	<i>ISO</i>
120	9.77	0.98	<i>ISO</i>
150	7.12	0.67	<i>ISO</i>
170	8.18	0.64	<i>ISO</i>
200	3.55	1.15	<i>ISO</i>
450	0.595	0.035	Holland et al. (2002)
850	0.097	0.005	Holland et al. (2002)

the asymmetry is significantly above the noise at both wavelengths, and there is a low probability of contamination from a background source.

In the modelling presented here, we use the model presented in Holland et al. (2002) for the smooth ring. This has a radial distribution of cross-sectional area  $\bar{\sigma}(r)$  – where  $\bar{\sigma}(r) dr$  is the fraction of the total cross-sectional area  $\sigma_{\text{tot}}$  between  $r$  and  $r + dr$  – that is indicative of a ring that has a mean radial distance of  $\sim 150$  au and a width of  $\sim 50$  au. This model is, however, described by the distribution of the orbital elements of the material in the ring (see e.g. WDT99). Material in their model has average orbital inclinations



**Figure 1.** The SED of emission from Fomalhaut after subtraction of the stellar photosphere (see Table 1). The asterisks, diamonds and triangles correspond to the *IRAS*, *ISO* and SCUBA fluxes, respectively. The lines represent model fits to these data (see Section 4) that assume that the disc’s spatial distribution is defined by a fit to the  $450\ \mu\text{m}$  image (Holland et al. 2002), and that it is comprised of solid grains of  $1/3$  (by volume) amorphous silicate and  $2/3$  organic refractory material (e.g. Li & Greenberg 1997, 1998). The grains’ size distributions in the models extend down to the radiation pressure blow-out limit,  $D_{\text{min}} = 7\ \mu\text{m}$ , with power law indices given by  $q_d$ .

and eccentricities of  $I = 5^\circ$  and  $e = 0.065$ . While this modelling could not uniquely constrain these parameters, these values result in a best fit to the data and are consistent with indications given later in this paper that we are observing a collisional cascade for which such high values are likely (Section 5.3).

### 3 THEORETICAL PLANETESIMAL SIZE DISTRIBUTION

#### 3.1 Basic idea

Debris discs are thought to have formed in the first 10 Myr or so of a star’s life through canonical planet formation processes, wherein collisions between initially submicron dust particles resulted in their coagulation and eventual growth into planetesimals and maybe even planets (e.g. Lissauer 1993). Toward the end of this process the collisional processing was reversed, whereupon collisions between planetesimals more frequently resulted in their net destruction rather than growth. Such a change is likely to have been the consequence of an increase in the average relative velocity of collisions, probably caused by some combination of the removal of the gas disc (which was acting to damp the eccentricities and inclinations of disc material) and gravitational perturbations from planets or large planetesimals that formed in the disc (Kenyon & Bromley 2001), although this cross-over stage is not well understood. In the debris disc systems, we know that by the end of this process, regardless of whether planets or planetesimals formed close to the star, a disc of planetesimals was able to form and subsequently survive in a ring at  $>30$  au from the star (Lagrange et al. 2001). Once their collisional velocities are pumped up, these planetesimals collide and break up into smaller fragments, initiating a collisional cascade. A typical planetesimal in such a cascade is most likely to have been created by the break-up of a larger parent body in a collision with another large planetesimal. This parent body would most likely have been created by the break-up of an even larger body, and the planetesimal itself will most likely end up as a parent body for planetesimals smaller than itself. It is the large quantities of small dust in the cascade, and its commensurate large surface area, which makes a debris disc observable. This dust is replenished as long as there remains a supply of large primordial bodies to feed the cascade.

#### 3.2 Size distribution of an infinite collisional cascade

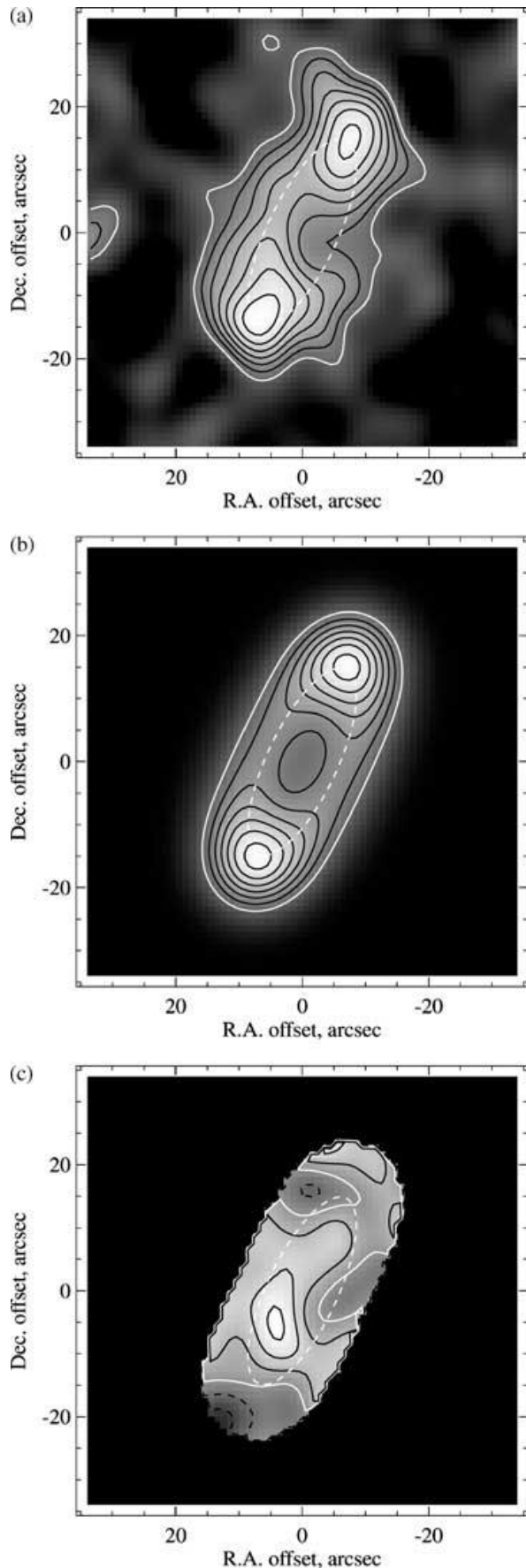
The equilibrium size distribution resulting from a collisional cascade can be found from theoretical arguments (Dohnanyi 1969):

$$n(D) \propto D^{2-3q_d}, \quad (1)$$

where  $n(D)dD$  is the number of planetesimals of size between  $D$  and  $D + dD$ . For a self-similar collisional cascade it can be shown that  $q_d = 1.833$  (Tanaka, Inaba & Nakazawa 1996), however if the strength of a planetesimal is size-dependent (see Section 5), then the slope of the equilibrium distribution could be slightly shifted from 1.833 (see Fig. 12 of Durda & Dermott 1997, hereafter DD97).

#### 3.3 Size distribution of small grains

The collisional cascade distribution is only expected to hold for disc particles that are large enough not to be affected by radiation forces (WDT99). Particles smaller than a few  $\mu\text{m}$  are blown out of the system by radiation pressure on orbital time-scales and all particles spiral into the star due to Poynting–Robertson (P–R) drag on time-scales that are proportional to their size. It can be shown, however, that the high densities of the observed extrasolar discs result in



collisional lifetimes for dust of all sizes that are shorter than their P–R drag lifetimes, and so the latter effect can be ignored (WDT99). Thus here we make the assumption that the size distribution holds down to the radiation pressure blow-out limit  $D_{\min}$  (i.e. those for which  $\beta = F_{\text{rad}}/F_{\text{grav}} > 0.5$ ), below which the distribution is cut-off.

A sharp cut-off below  $D_{\min}$  is, however, a simplification. For a start, a sharp cut-off would cause a ‘wave’ in the equilibrium size distribution (Durda, Greenberg & Jedicke 1998, hereafter DGJ98), because in an equilibrium situation particles close to the blow-out limit should have been broken up by the missing population of particles smaller than themselves. As these are not removed, their number increases and this affects the particles they are expected to break up, and so on. Also, while it may be possible to ignore the contribution of particles smaller than the blow-out limit to the total cross-sectional area of disc material (e.g. if their lifetime is much shorter than their collisional lifetime; WDT99), the size distribution of gravitationally bound grains close to the blow-out limit could be affected by their destruction in collisions with the blow-out grains (Krivov, Mann & Krivova 2000). The size distribution of these bound grains would be further complicated by radiation pressure which causes them to have large eccentricities, thus reducing their number density in the disc region (see e.g. Fig. 4 of Augereau et al. 2001).

### 3.4 Size distribution of large planetesimals

At the large-size end, planetesimals that have a collisional lifetime that is longer than the age of the system do not form part of the collisional cascade, and their distribution must be primordial. However, it is unclear what this would be, as even the primordial distribution of solar system objects is unknown. It is uncertain whether the distribution of asteroids larger than 30 km is primordial due to uncertainties in the collisional properties of these asteroids (DGJ98). The distribution of Kuiper belt objects larger than 100 km, which appear to have a relatively steep distribution ( $q_d = 1.83\text{--}2.27$ , Luu & Jewitt 1998; Gladman et al. 1998), is also difficult to interpret, because while the objects themselves might be primordial, their distribution might not be. The sweeping of Neptune’s resonances through the young Kuiper belt as Neptune’s orbit expanded early in the history of the solar system (Hahn & Malhotra 1999) would have resulted in a size distribution of objects trapped in resonance that is steeper than the primordial distribution (Jewitt 1999). The primordial distribution expected from theoretical grounds is also unclear, as simulations of planet formation result in distributions that are usually steeper than the cascade distribution, but with  $q_d$  anywhere between 2 and 2.5 (e.g. Kokubo & Ida 2000; Kenyon 2002). In any case, the size distribution of the largest planetesimals in the

**Figure 2.** 450- $\mu\text{m}$  images of the Fomalhaut disc (adapted from Holland et al. 2002): (a) observation taken with SCUBA at the JCMT, (b) axisymmetric smooth disc model, and (c) the residuals (= observation – model), which show that the asymmetry in the observation could be explained by a bright clump embedded in the smooth disc. All contours are spaced at  $1\sigma = 13 \text{ mJy beam}^{-1}$  levels. For (a) and (b) the lowest (solid white) contour is set at  $3\sigma$ , while for (c) the solid white line indicates the zero level and the black solid and dashed lines are the positive and negative  $1\sigma$  levels, respectively. The dashed white oval shows the approximate inner edge of the mid-plane of the disc, a 125-au radius ring inclined at  $20^\circ$  to the line of sight. The coordinates are given as their offsets from the position of the star. The stellar photosphere has been subtracted from the observation.

distribution is likely to be dominated by the runaway growth process, whereby these grew faster than the rest of the population, causing a bimodal size distribution (e.g. Wetherill & Stewart 1993; Kokubo & Ida 1996). Runaway growth is discussed further in Section 7.2.

## 4 SED MODELLING

### 4.1 The SED model

The SED of a disc is determined by three factors: the spatial distribution of material in the disc, its size distribution, and its optical properties (determined by its composition). While these factors could be interrelated, we assume here that they are independent. For the spatial distribution we assume that particles of all sizes have the spatial distribution inferred from the modelling of the 450- $\mu\text{m}$  images defined by  $\bar{\sigma}(r)$  (see Section 2). For the size distribution we consider one that follows equation (1), and assume that this distribution is cut-off below  $D_{\min}$ , the radiation pressure blow-out limit, and extends up to sizes large enough that this limit does not affect the SED.

For the composition of the particles we use the core–mantle model developed by Li & Greenberg (1997) for interstellar dust, which has since been applied to cometary dust (Greenberg 1998) and to dust grains in extrasolar discs (Li & Greenberg 1998; Augereau et al. 1999). It seems reasonable to apply this model to extrasolar discs, as these grains are likely to be reprocessed (albeit significantly) interstellar grains. In this model the dust is assumed to be aggregates of core–mantle grains with a silicate core ( $\rho = 3500 \text{ kg m}^{-3}$ ) and an organic refractory mantle comprised of ultraviolet (UV) photoprocessed ices which accreted on to the silicate cores in the interstellar medium ( $\rho = 1800 \text{ kg m}^{-3}$ ). These aggregates are assumed to be porous and to have a fraction of water ice ( $\rho = 1200 \text{ kg m}^{-3}$ ) filling the gaps; this is ice that would have frozen on to the growing grains in the protoplanetary disc. As interstellar grains are amorphous and would only be crystallized at the high temperatures expected close to stars, we assume that both the silicate and the ice components of the grains are amorphous. We also assume a volume fraction of silicates to organic refractory material in the core–mantle of 1:2 as found for interstellar-type grains (Li & Greenberg 1997). Thus the particles' composition is defined by their porosity,  $p$ , and the volume fraction of the gap in the grains that is filled with ice,  $q_{\text{ice}}$ . We use the optical constants for each of the components taken from Li & Greenberg (1997, 1998) and compute those of the composite material using Maxwell–Garnett Effective Medium theory (Bohren & Huffman 1983). The absorption efficiencies of these grains,  $Q_{\text{abs}}(\lambda, D)$ , are then calculated using Mie theory (Bohren & Huffman 1983), Rayleigh–Gans theory or geometric optics in the appropriate limits (see Laor & Draine 1993). These then define the temperature of the grains,  $T(D, r)$ , and their radiation pressure blow-out limit,  $D_{\min}$  (see e.g. WDT99).

Thus the SED model is defined by the following four parameters:  $\sigma_{\text{tot}}$  (the total cross-sectional area in the disc),  $q_d$ ,  $p$ , and  $q_{\text{ice}}$ . The flux in  $J_\nu$  from the disc at a given wavelength is given by (e.g. WDT99)

$$F_\nu = 2.35 \times 10^{-11} R_\star^{-2} \sigma_{\text{tot}} \int_{D_{\min}}^{D_{\max}} Q_{\text{abs}}(\lambda, D) \bar{\sigma}(D) \times \int_{r_{\min}}^{r_{\max}} B_\nu[T(D, r)] \bar{\sigma}(r) dr dD, \quad (2)$$

where  $\bar{\sigma}(D) = 0.25\pi D^2 n(D) / \sigma_{\text{tot}}$  and  $\sigma_{\text{tot}}$  is in  $\text{au}^2$ . Here we have assumed that the disc is optically thin at all wavelengths and that there is no disc self-heating.

### 4.2 The modelling procedure

We determined the best-fitting  $q_d$  for each particle composition (defined by  $p$  and  $q_{\text{ice}}$ ) by minimizing  $\chi^2 = \sum_{i=1}^{N_{\text{obs}}} [(F_{\nu_i, \text{obs}} - F_{\nu_i, \text{mod}}) / dF_{\nu_i, \text{obs}}]^2$ , where  $\sigma_{\text{tot}}$  was set to scale the SED to fit the 850  $\mu\text{m}$  flux. Fig. 1 shows how this procedure was applied for solid grains (for which  $p = 0$ ,  $D_{\min} = 7 \mu\text{m}$  and  $\rho = 2370 \text{ kg m}^{-3}$ , and which by definition contain no water ice from the protoplanetary disc). The best fit gave  $q_d = 1.84$ ,  $\sigma_{\text{tot}} = 33.7 \text{ au}^2$  and  $\chi^2 = 43.4$ . This type of modelling constrains  $q_d$  very well because it has a large effect on the SED: increasing  $q_d$  increases the contribution of small grains and so results in an increase in the flux shortward of about 60  $\mu\text{m}$  (due to the higher temperatures of small grains) but a decrease in the flux at longer wavelengths (as small grains emit less efficiently than larger ones at long wavelengths).<sup>1</sup>

Increasing  $q_d$  also steepens the submm spectral slope, defined by  $\alpha$ , where  $F_\nu \propto \lambda^{-\alpha}$ , because the emission here is dominated by grains of a size comparable with the wavelength so that the ratio of fluxes at two wavelengths is determined by the ratio of cross-sectional area in grains of size comparable to those wavelengths. This effect can be quantified by a simple model in which all the dust is at the same distance from the star and has the same temperature,  $T$ , and has emission efficiencies that are given by

$$Q_{\text{abs}}(\lambda, D) = \begin{cases} 1 & \text{for } \lambda < D \\ (\lambda/D)^n & \text{for } \lambda > D. \end{cases} \quad (3)$$

Putting this into equation (2), we find that in the Rayleigh–Jeans regime:

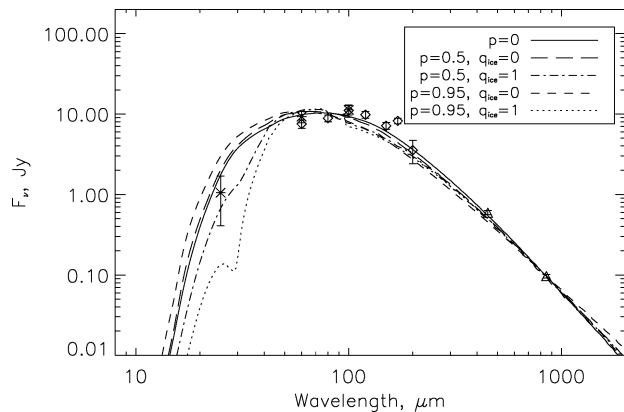
$$F_\nu \propto \begin{cases} \lambda^{3-3q_d} & \text{for } n > 3q_d - 5 \\ \lambda^{-2-n} & \text{for } n < 3q_d - 5, \end{cases} \quad (4)$$

i.e. increasing  $q_d$  steepens the slope up to a constant value for  $q_d > 5/3 + n$ . Thus the submm spectral slope is largely determined by the size distribution of particles of submm sizes. This demonstrates the importance of an SED model not only giving a good  $\chi^2$  fit to the whole SED, but also giving a good fit to  $\alpha_{\text{obs}} = 2.85$ . In this simple model, the observed slope implies that either  $q_d = 1.95$  and  $n > 0.85$ , or  $n = 0.85$  and  $q_d > 1.95$ . In fact, we find that for solid grains  $\alpha - \alpha_{\text{obs}} \approx 4.9(q_d - 1.86)$  and so  $q_d = 1.84$  fits the submm slope relatively well. The discrepancy from the simple model is because real grains have emission features that are not reproduced by equation (3).

### 4.3 Modelling results

The analysis in Fig. 1 was repeated using grains with a range of porosities and ice fractions, but found that increasing either  $p$  or  $q_{\text{ice}}$  resulted in a worse fit to the data (i.e. a higher  $\chi^2$ , see Fig. 3). The problem with increasing the porosity of the grains is that even though solid grains are less affected by radiation pressure than porous grains are (mostly due to their higher density), and so can exist in the disc down to much smaller sizes (see  $D_{\min}$  in the caption of Fig. 3), the smallest porous grains of the disc are still hotter than its smallest solid grains because they appear to be made up of an agglomeration of much smaller particles. Thus the SED resulting from a porous-grained disc peaks at shorter wavelengths than a solid-grained one (see Fig. 3). Higher porosity grains also have fewer emission features, which results in a flatter submm slope. The

<sup>1</sup> It is not immediately obvious from Fig. 1 that this cross-over occurs at 60  $\mu\text{m}$  because, as these models have been scaled to  $F_\nu(850 \mu\text{m})$ , those with a higher  $q_d$  also have a higher  $\sigma_{\text{tot}}$ .

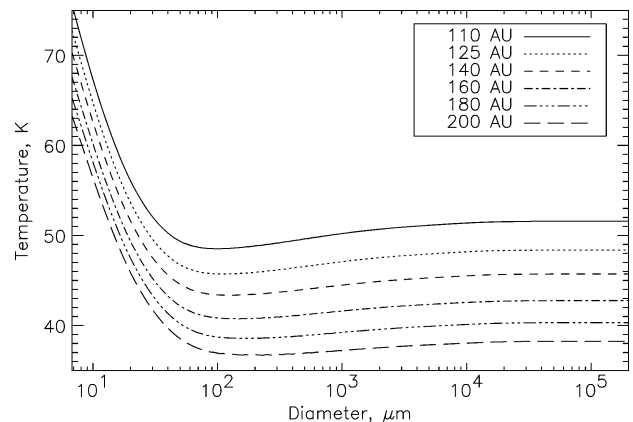


**Figure 3.** The best fits to the observed SED for five models with different grain compositions (see legend). The resulting best fits for the size distribution, the total cross-sectional area in the disc, and the  $\chi^2$  for the five models (top to bottom in the legend) are:  $q_d = [1.84, 1.81, 1.87, 1.81, 1.88]$ ,  $\sigma_{\text{tot}} = [33.7, 33.2, 27.1, 34.7, 23.4]$  au<sup>2</sup>, and  $\chi^2 = [43.4, 75.3, 101, 118, 117]$ . The minimum grain size in these models is:  $D_{\text{min}} = [7, 13, 9, 124, 11]$   $\mu\text{m}$ .

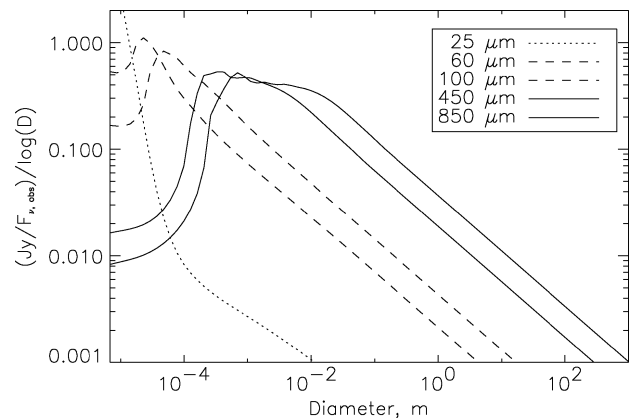
problem with increasing the ice fraction in the grains is that they have a range of emission features in the far-IR, but not of a shape that increases the submm slope to the same levels as for non-icy grains. The resulting spectrum is too peaked in the 40–90  $\mu\text{m}$  wavelength region (see Fig. 3), and because icy grains are inefficient absorbers of starlight, their cool temperature results in a poor fit to the mid-IR flux. A poor fit to the 25- $\mu\text{m}$  flux would not in itself rule out a model, as there may be a contribution to the observed flux from an, as yet undetected, hotter disc component such as that inferred to exist in the HR 4796 system (Augereau et al. 1999). We conclude that while solid non-icy grains are the most likely candidates for the grains in this disc, SED modelling cannot set strict constraints on the composition of the grains without a more accurate determination of the mid-IR emission of the disc. However, a more robust feature of the SED modelling is the size distribution parameter  $q_d$ , which lies in the range 1.81–1.88 for all models.

In the rest of this paper we use the parameters ( $D_{\text{min}}$ ,  $\rho$ ,  $q_d$  and  $\sigma_{\text{tot}}$ ) derived from this modelling assuming solid non-icy grains, as this composition is also consistent with that expected on theoretical grounds. Grains that grow through grain–grain collisions in a protoplanetary disc are expected to have significant porosity (e.g. Wurm & Blum 1998), similar to that of grains resulting from the sublimation of comets in the solar system ( $p = 0.85$ – $0.95$ ). Grain–grain collisions, if energetic enough however, can lead to significant grain compaction (Dominik & Tielens 1997). Thus we expect grains produced by comet sublimation to have a high, primordial-like, porosity (such as has been inferred for the grains around around  $\beta$  Pictoris and HR 4796A; Li & Greenberg 1998; Augereau et al. 1999), whereas grains at the end of a collisional cascade should be more compact, maybe with a porosity close to that of stone meteorites ( $p = 0.05$ – $0.3$ ; Flynn, Moore & Klöck 1999). The size distribution we have inferred,  $q_d$  close to 1.84, is consistent with the dust having its origin in a collisional cascade (Section 3.2); collisional lifetime arguments in the following section also support this hypothesis. Also, we expect little contribution to the dust population from cometary sublimation because of the low temperatures in the disc (see Fig. 4). Thus a low level of porosity is to be anticipated for the dust.

It is harder to predict the amount of water ice on the grains. Certainly the low temperatures in the disc imply that while the grains



**Figure 4.** The temperatures of different sized dust grains at different distances from Fomalhaut, assuming these are solid grains composed of 1/3 (by volume) amorphous silicate and 2/3 organic refractory material.



**Figure 5.** The contribution of different sized dust grains to the observed fluxes for the best fit model to the SED (i.e. that shown in Fig. 1 with  $q_d = 1.84$ ). The area under the curves between two grain sizes defines the proportion of the flux that originates from this size range; the total area across all sizes is unity.

were forming (i.e. before they were compacted by collisions), water vapour should have condensed on to the aggregates, at least partially filling the gaps in the grains (Pollack et al. 1994; Li & Greenberg 1998). However, this ice may have since been sublimated by repeated collisional heating, removed by sputtering by stellar wind ions (Jurac, Johnson & Donn 1998). Or indeed this ice may still exist on the grains, yet it could look more like the organic refractory material in our model due to the formation of an irradiation mantle by stellar UV photons and galactic cosmic rays (e.g. Johnson et al. 1987; Cooper, Christian & Johnson 1998; Jewitt 1999). We note that water ice features have been detected in some, but not all, Kuiper belt objects (e.g. Luu, Jewitt & Trujillo 2000; Brown, Blake & Kessler 2000).

#### 4.4 Limitations for large planetesimals

The size distribution inferred from the SED modelling can only be applied to grains that contribute to the observed SED fluxes. Fig. 5 shows that planetesimals larger than  $\sim 0.2$  m in diameter contribute less than 5 per cent of the 850  $\mu\text{m}$  flux (and even less at shorter wavelengths). Thus we can only derive the size distribution for grains that are smaller than  $\sim 0.2$  m (although, as we will see in

the next section, we can extend this up to larger sizes by collisional lifetime arguments). Because the total mass of the Fomalhaut disc, assuming that the size distribution with  $q_d = 1.84$  extends up to  $D_{\max}$ , is given by

$$m_{\text{tot}}/M_{\oplus} \approx 0.5D_{\max}^{0.48}, \quad (5)$$

where  $D_{\max}$  is in m, this means that we see a dust mass of  $\sim 0.2M_{\oplus}$ .

## 5 CATASTROPHIC COLLISIONS

The outcome for a planetesimal of size  $D$  when it is impacted by another of size  $D_{\text{im}}$  at a velocity  $v_{\text{col}}$  is determined by the specific incident kinetic energy,

$$Q = 0.5(D_{\text{im}}/D)^3 v_{\text{col}}^2. \quad (6)$$

The shattering threshold,  $Q_S^*$ , is defined as the specific incident energy required to break up the planetesimal, resulting in fragments with a range of sizes up to half the mass of the original planetesimal. Collisions with  $Q < Q_S^*$  result in cratering whereby some material is ejected, but the planetesimal remains largely intact. More energetic collisions result in the planetesimal being reduced to more and smaller fragments.<sup>2</sup> In general, larger planetesimals have lower shattering thresholds, as they contain larger flaws which are activated at lower tensile stresses (Housen & Holsapple 1990, 1999). Very large planetesimals, however, are strengthened by their gravitational self-compression which inhibits crack propagation (Davis et al. 1985). After the collision, the fragments disperse as a result of the kinetic energy imparted to them in the collision. For the break-up of a large planetesimal ( $D > 150$  m), however, this energy may not be enough to overcome the gravitational binding energy and some fragments may reaccumulate into a rubble pile (e.g. Campo Bagatin, Petit & Farinella 2001; Michel et al. 2001). A catastrophic collision is defined as one in which  $Q > Q_D^*$ , where the dispersal threshold,  $Q_D^*$ , refers to an impact in which the largest fragment resulting from the collision after reaccumulation has taken place (i.e. this fragment could be a rubble pile) has half the mass of the original planetesimal. The break-up of a small planetesimal (i.e.  $D < 150$  m), for which  $Q_D^* \approx Q_S^*$ , is said to occur in the strength regime, while the break-up of a larger planetesimal is said to occur in the gravity regime, where  $Q_D^* \gg Q_S^*$ .

The catastrophic collisional lifetime of a planetesimal depends on the size distribution of the disc, as only collisions in which  $D_{\text{im}} > D_{\text{ic}} = X_{\text{ic}}D$  have enough energy to be destructive, where

$$X_{\text{ic}} = (2Q_D^*/v_{\text{ic}}^2)^{1/3}. \quad (7)$$

We use the term ‘threshold–catastrophic’ to denote a collision between planetesimals of size  $D$  and  $D_{\text{ic}}(D)$ , and  $v_{\text{ic}}(D)$  is the relative velocity of such collisions. Here we assume that all collisions have the same average relative velocity at large separations,<sup>3</sup>

$$v_{\text{rel}} = f(e, I)v_k, \quad (8)$$

<sup>2</sup>For collisions with very small relative velocities there are two other possible collisional outcomes: the planetesimals could rebound in an inelastic collision, or they could stick together. While these possibilities are important in protostellar discs, they are not considered in this paper, as the collisional velocities in debris discs are thought to be too high for these modes to be important.

<sup>3</sup>The distributions of the eccentricities and inclinations of planetesimals predicted from planet formation models are dependent on their size. However, in a debris disc undergoing a collisional cascade these distributions would depend on the mechanism which causes the high eccentricities and inclinations in these discs; e.g. if this mechanism is the continual gravitational stirring of massive planetesimals or planets in the disc, these distributions

where  $f(e, I)$  is a function of the average eccentricities and inclinations of the planetesimals given as  $\sqrt{1.25e^2 + I^2}$  (Lissauer & Stewart 1993; Wetherill & Stewart 1993), and  $v_k$  is the Keplerian velocity at this distance from the star. In the Fomalhaut model  $f(e, I) \approx 0.11$ ; thus  $v_{\text{rel}} \approx 0.4 \text{ km s}^{-1}$  at the mean distance of 150 au. The relative velocity of a collision is then

$$v_{\text{col}}^2 = v_{\text{rel}}^2 + v_{\text{esc}}^2(D, D_{\text{im}}), \quad (9)$$

where the mutual escape velocity of the two planetesimals is given by

$$v_{\text{esc}}^2(D, D_{\text{im}}) = (2/3)\pi G\rho \frac{D^3 + D_{\text{im}}^3}{D + D_{\text{im}}}. \quad (10)$$

The increase in impact velocity due to gravity (i.e. gravitational focusing) becomes important for collisions for which  $v_{\text{esc}} > v_{\text{rel}}$ , i.e. for planetesimals with  $D > v_{\text{rel}}/\sqrt{(2/3)\pi G\rho}$ . For Fomalhaut this corresponds to planetesimals larger than about 700 km in diameter, or 0.6 per cent of a lunar mass.

### 5.1 Dispersal threshold model

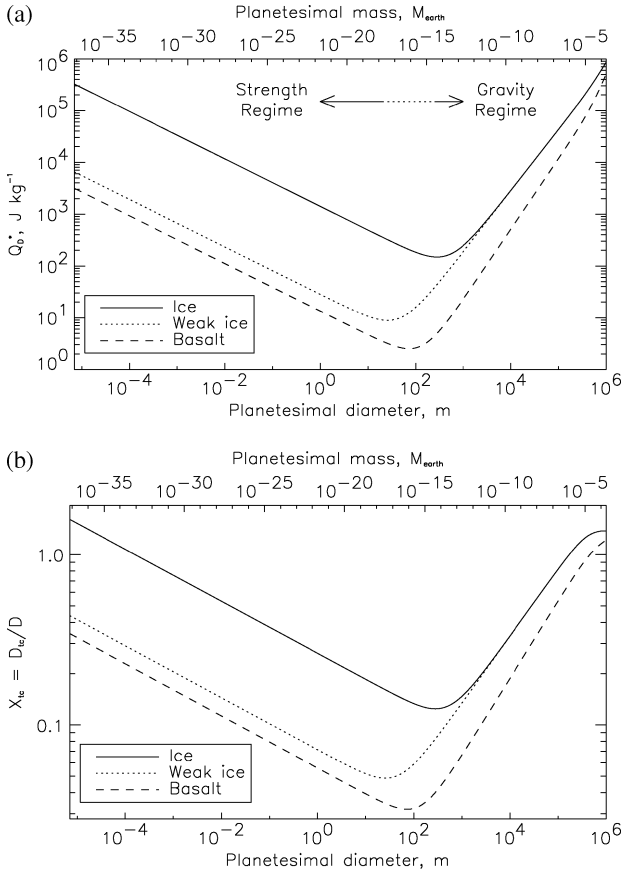
Several authors have studied how  $Q_S^*$  and  $Q_D^*$  vary with planetesimal diameter for a variety of materials (e.g. rocks/ice) with a range of structural properties (e.g. homogeneous/rubble pile), and for collisions with a range of relative velocities (e.g.  $\text{km s}^{-1}$  debris disc-like/m  $\text{s}^{-1}$  protostellar nebula disc-like impacts) and impact parameters (e.g. head-on/glancing blow). The techniques that have been used to study collisions range from laboratory experiments which provide a direct measure of a collisional outcome (e.g. Fujiwara et al. 1989; Housen & Holsapple 1999), to theoretical studies (e.g. Petit & Farinella 1993), to interpretation of the distributions of the asteroid families (e.g. Tanga et al. 1999; Cellino et al. 1999) or the main-belt population (DD97; DGJ98), to computational modelling using smooth particle hydrodynamics (SPH; Love & Ahrens 1996; Benz & Asphaug 1999, hereafter BA99). However, most of these studies are valid only in specific regimes (e.g. in the strength or gravity regimes).

In this paper we use the results of BA99, who use SPH to model the fragmentation of both solid ice and basalt of a broad range of sizes (up to 200 km in diameter) for impacts with a range of impact parameters and with relative velocities comparable to those expected in extrasolar planetesimal discs. In the rest of this paper  $Q_D^*$  refers to the dispersal threshold averaged over all impact parameters (BA99). The BA99 models reproduce the well-known features of the  $Q_D^*$  versus  $D$  plot; i.e. the dispersal threshold decreases with size [ $\propto D^{-(0.36-0.45)}$ ] in the strength regime as a result of the lower shattering strength of larger planetesimals, but increases with size [ $\propto D^{1.19-1.36}$ ] in the gravity regime as a result of the extra energy required to impart enough kinetic energy to the collisional fragments to overcome the gravity of the planetesimal. However, the impact strengths they derive for cm-sized ice grains impacted at  $\sim 3 \text{ km s}^{-1}$  are 20–50 times higher than those directly measured in the laboratory (e.g. Arakawa 1999; Ryan, Davis & Gliblin 1999). Their results for basalt impacted at  $3 \text{ km s}^{-1}$  do not show the same discrepancy (Holsapple 1994).

While the SED modelling suggested that there is little water ice in the Fomalhaut grains, the structural properties of the organic

would be uniform with planetesimal size, whereas if this mechanism was stirring by a passing star these distributions would evolve with time after the event and would be size-dependent (Kenyon & Bromley 2002).

refractory mantle would be close to that of the ice studied in BA99, as the mantle is essentially comprised of UV photoprocessed ices. Thus the collisional properties of planetesimals in Fomalhaut’s disc would lie somewhere between those of rock and ice. Here we consider three different models for  $Q_D^*$ : Ice refers to the BA99 result for ice impacted at 0.5 and 3 km s<sup>-1</sup>; Weak Ice also refers to the BA99 results for ice, except that the part of a planetesimal’s  $Q_D^*$  resulting from its strength is reduced by a factor of 50 to reflect the discrepancy with laboratory results; and Basalt refers to the BA99 result for basalt impacted at 3 and 5 km s<sup>-1</sup>. In all cases we have used  $\rho = 2370 \text{ kg m}^{-3}$  and the BA99 results at the two impact velocities have been scaled to the relative velocity of a threshold–catastrophic collision,  $v_{tc}$ , assuming that  $Q_D^* \propto v_{col}^{\gamma(D)}$ . This procedure is iterative, because to calculate  $Q_D^*$  we need to know  $v_{tc}$ , for which we need to know  $X_{tc}$ , which is itself determined by both  $Q_D^*$  and  $v_{tc}$ . The resulting models for  $Q_D^*$  and  $X_{tc}$  are shown in Fig. 6.



**Figure 6.** The collisional properties of planetesimals of different sizes in the Fomalhaut disc for the three models (Ice, Weak Ice, and Basalt) based on the SPH modelling of Benz & Asphaug 1999: (a) the dispersal threshold; (b) the relative size of an impactor required to catastrophically break up a planetesimal. The dispersal threshold decreases with size in the strength regime ( $D < 150 \text{ m}$ ) as a result of the lower shattering strength of larger planetesimals, but increases with size in the gravity regime due to the extra energy required to impart enough kinetic energy to the collisional fragments to overcome the planetesimal’s gravity. The relative size required for a catastrophic impact varies with  $D$  as does  $Q_D^*$  except that it flattens off for  $D > 700 \text{ km}$  because the extra energy required to disperse the collisional fragments is provided by the higher impact velocities, which have been enhanced by the planetesimals’ gravitational attraction (from  $\sim 0.4 \text{ km s}^{-1}$  at  $100 \text{ km}$  to  $\sim 0.8 \text{ km s}^{-1}$  at  $1000 \text{ km}$ ).

While the BA99 models apply only to planetesimals up to 200 km in diameter, we have assumed that their results can also be scaled up to 1000 km in diameter. We do not expect their models to be applicable much beyond this, as in this regime collisions occur at the planetesimals’ escape velocity. As some energy is lost in the collision, only a small fraction of the collisional fragments can escape, and so such collisions must result in net accretion rather than destruction (i.e.  $Q_D^* \rightarrow \infty$ ). Such an effect would not have been picked up in BA99, as their experiments were all carried out with impact velocities that are more than an order of magnitude under the escape velocities. Collisions between such massive bodies, which we further expect to be rubble piles, are discussed in Section 7.

## 5.2 Catastrophic collision time-scale

The rate of impacts from planetesimals in the size range  $D_{im}$  to  $D_{im} + dD_{im}$  falling on to a planetesimal of size  $D$  is given by  $R_{col}(D, D_{im})dD_{im}$ , where (e.g. Opik 1951)

$$R_{col}(D, D_{im}) = f(D, D_{im})\sigma(r, \theta, \phi)v_{rel}. \quad (11)$$

In this expression,  $\sigma(r, \theta, \phi)$  is the average cross-sectional area density of disc material, which, for a ring of radius  $r$  and radial width  $dr$ , and that has a (small) mean orbital inclination of particles in the ring of  $I$ , is approximately

$$\sigma(r, \theta, \phi) = \sigma_{tot}/(4\pi r^2 dr I). \quad (12)$$

Also,  $f(D, D_{im})\sigma(r, \theta, \phi)dD_{im}$  is the cross-sectional area density of planetesimals in the given size range that the planetesimal of size  $D$  sees, where

$$f(D, D_{im}) = \bar{\sigma}(D_{im}) \left( \frac{D + D_{im}}{D_{im}} \right)^2 \left[ 1 + \frac{v_{esc}^2(D, D_{im})}{v_{rel}^2} \right]. \quad (13)$$

The last term in equation (13) is the gravitational focusing factor, which becomes important for collisions with planetesimals in the Fomalhaut disc that are larger than about 700 km.

The catastrophic collision time-scale of a planetesimal is thus given by

$$t_{cc}(D) = t_{per}(r dr/\sigma_{tot})[2I/f(e, I)]/f_{cc}(D), \quad (14)$$

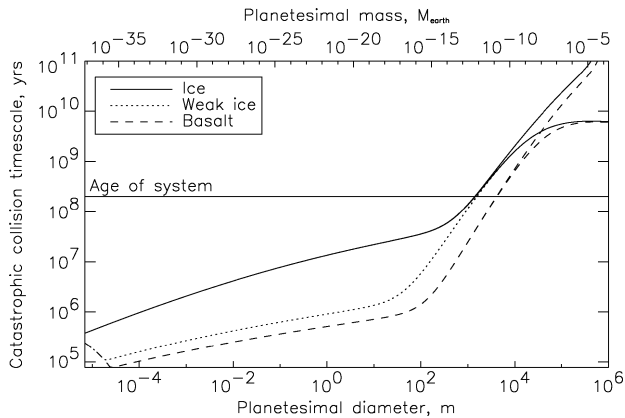
where  $t_{per}$  is the orbital period at  $r$ , and  $f_{cc}(D)$  is defined by

$$f_{cc}(D) = \int_{D_{tc}(D)}^{D_{max}} f(D, D_{im})dD_{im}, \quad (15)$$

where  $D_{tc}$  is the larger of  $X_{tc}D$  or  $D_{min}$ .

Fig. 7 shows the catastrophic collisional lifetime of planetesimals of different sizes in the Fomalhaut disc for the three models of their collisional properties, using the size distribution inferred from the SED modelling ( $q_d = 1.84$ ), assuming that this is truncated at 1000 and 10 000 km. Normally a planetesimal is broken up by one that just has enough energy to do so (i.e. in a threshold–catastrophic collision), because out of all the planetesimals that can destroy it, these have the largest total cross-sectional area in the disc. However, the enhanced cross-sectional area of gravitationally focusing planetesimals can mean that a planetesimal is more likely to end its life by being accreted on to one of these than by being destroyed by a smaller planetesimal, hence the shorter collisional lifetime of large planetesimals when  $D_{max} = 10\,000 \text{ km}$ . As there is less cross-sectional area available to destroy larger planetesimals, there is a steady increase in collisional lifetime with size; e.g. a simple model which assumes that most of this area is in planetesimals of size  $X_{tc}D$ , where  $X_{tc} < 1$ , shows that if  $Q_D^* \propto D^a$  then  $t_{cc} \propto D^{(3+a)q_d - (5+a)}$ , implying that the BA99 models should have  $t_{cc} \propto D^{0.14-0.22}$  in the strength regime and  $t_{cc} \propto D^{1.52-1.66}$  in the gravity regime, consistent with Fig. 7.





**Figure 7.** The collisional lifetime of planetesimals of different sizes in the Fomalhaut disc. The solid, dotted and dashed lines are for the three models for the collisional properties of these planetesimals described in the text. The two lines for each model assume that the size distribution inferred from the SED modelling ( $q_d = 1.84$ ) extends up to planetesimals of size 1000 and 10 000 km (distributions with larger planetesimals have lower collisional lifetimes).

### 5.3 Implications for Fomalhaut’s size distribution

While the collisional lifetime of all planetesimals contributing to the Fomalhaut observations ( $D < 0.2$  m) is model-dependent, because each requires a collision with different sized planetesimals for catastrophic destruction (Fig. 6b), all models predict a lifetime that is shorter than the age of the system (200 Myr) by more than an order of magnitude (Fig. 7). Thus the primordial populations of this size range of particles has already been substantially depleted. However, these populations will also have been replenished with collisional fragments from the break-up of larger planetesimals. As the largest particles that we can see are replenishing the smaller particles, it is an unavoidable conclusion that this part of the size distribution forms a collisional cascade. The real questions are: (i) is the cascade in equilibrium, and (ii) how far does it extend up in size, or in other words, what is feeding the cascade?

In answer to (i), while the parametrization of the size distribution assumed in the SED modelling is only an approximation (e.g. Sections 3.2 and 3.3), the derived slope,  $q_d = 1.84$ , is very close to that expected for an equilibrium cascade. Also, the short collisional lifetimes in Fig. 7 imply that the cascade should have reached equilibrium within the age of the system.

In answer to (ii), the short collisional lifetime of the 0.2-m diameter planetesimals we see today (0.4–9 Myr depending on their composition, see Fig. 7) implies that if there were no larger planetesimals with which to replenish their number, then while all the 0.2-m planetesimals we see must be primordial, their current population should only be a small fraction of their original population. In fact, it must be an infinitely small fraction of the original population, as the collisional lifetime of these planetesimals would have been even shorter at earlier epochs when their population was more massive (this lifetime is approximately inversely proportional to its mass). In other words, unless the collisional cascade was only initiated in the last few Myr, the large number of 0.2-m planetesimals we observe today implies that there must be a population of larger bodies. Similar arguments can be applied for all planetesimals for which their collisional lifetime is shorter than the time since collisions became destructive (which we assume here is approximately the age of the system); the frequency of collisions amongst these large bodies means that their distribution would also be expected to follow that

of a collisional cascade. Thus we infer that the collisional cascade starts with planetesimals of a size for which their collisional lifetime is equal to the age of the system, i.e. those with  $D_{\max(\text{cc})} = 1.5\text{--}4$  km (Fig. 7), implying that the mass of material in the collisional cascade is  $m_{\text{cc}} \approx 20\text{--}30 M_{\oplus}$  (equation 5). If we had assumed that collisions became destructive more recently than 200 Myr ago, we would have inferred that the cascade started with smaller planetesimals. However, we note that a few km is also the size of planetesimal thought to have fed the collisional cascade in the young ( $< 1$  Gyr) Kuiper belt (Kenyon 2002).

We also predict that any planetesimals larger than 1.5–4 km are predominantly primordial and so their size distribution is also primordial; these will have contributed little to the collisional cascade. Their size distribution can be somewhat constrained from considerations of the total mass of the disc, which is dominated by that of the largest bodies. If the primordial size distribution connects smoothly to the collisional cascade distribution, and can be defined by the power-law exponent  $q_p$  between  $D_{\max(\text{cc})}$  and  $D_{\max}$ , then the total mass of primordial material would be

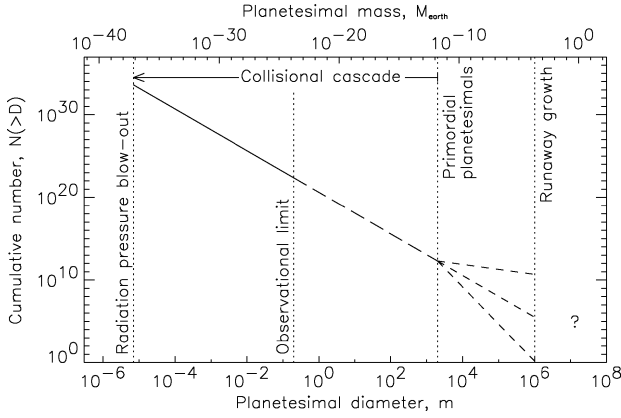
$$m_p/m_{\text{cc}} = 0.48(6 - 3q_p)^{-1} \left[ \left( D_{\max}/D_{\max(\text{cc})} \right)^{6-3q_p} - 1 \right]. \quad (16)$$

The minimum mass solar nebula (Weidenschilling 1977), if it extended out to such large distances, would have had a mass of solid material of 0.1–0.2  $M_J$  in the region 125–175 au. Planet formation models predict that for the gravitational perturbations of large planetesimals in the disc to have stirred the disc sufficiently to ignite a collisional cascade within 200 Myr, the planetesimals in Fomalhaut’s disc would have had to have grown to at least 1000 km in diameter and the disc must be at least 10 times more massive than the minimum mass solar nebula (see Fig. 5 of Kenyon & Bromley 2001). Thus a reasonable distribution for the primordial planetesimals, and one which we will use for reference later in the paper, is one with  $q_p = 1.84$  and  $D_{\max} = 1000$  km, as such a disc would have a mass of  $\sim 1.1 M_J$ , i.e. 5–10 times that of the minimum mass solar nebula. We note that 1000 km is similar in size both to the largest asteroids and Kuiper belt objects (Trujillo et al. 2001) as well as to the runaway growth limit of  $\sim 700$  km derived from the proposed eccentricities and inclinations of the orbits of Fomalhaut’s planetesimals (e.g. Section 5). The distribution of runaway planetesimals is discussed in Section 7.

We note that the shape of the primordial size distribution determines how the brightness of the disc evolves with time, and that this evolution can be constrained using the distribution of the brightnesses of observed debris discs. However we leave this discussion to a future paper. Our understanding of the size distribution of planetesimals in Fomalhaut’s disc from this study is summarized in Fig. 8.

## 6 INTRINSIC CLUMPINESS

Collisions between planetesimals cause the disc to be intrinsically clumpy, as after a collision the fragments of the target and the impactor move away from the impact site at a velocity determined by the kinetic energy imparted to them in the collision. This means that the debris from each planetesimal forms an expanding clump, the centre of which follows the orbit of the parent planetesimal. At any one time we would expect to see clumps in the disc with a range of both fluxes (those from different magnitude collisions) and sizes (smaller ones that were created recently and older, more extended, clumps). The probability of our witnessing a collisional clump of a given brightness and physical size depends on: the rate of



**Figure 8.** Inferred size distribution of planetesimals in Fomalhaut’s disc: a collisional cascade extends from  $\sim 4$  km planetesimals down to  $7 \mu\text{m}$  dust grains. Smaller dust grains are blown out of the system by radiation pressure. At submm and shorter wavelengths, we only see those members smaller than  $0.2$  m. Planetesimals larger than  $\sim 4$  km have an unknown primordial size distribution, however a reasonable distribution may be one that extends up to  $1000$  km with the same slope as the cascade (see Section 5.3). Planetesimals larger than  $\sim 1000$  km, if they exist, must have grown by runaway growth and would form a separate population (see Section 7.2).

collisions between different sized planetesimals (Section 6.1); the amount of dust such collisions produce, which depends on the size distribution of collisional fragments (Section 6.2); and on the rate at which this dust precesses around the orbit of the parent planetesimal, which depends on the ejection velocity of collisional fragments (Section 6.3).

### 6.1 Collision rates

The number of collisional events per unit time that occur in the disc between (target) planetesimals of size  $D$  to  $D + dD$  and those (impactors) of size between  $D_{\text{im}}$  and  $D_{\text{im}} + dD_{\text{im}}$  is  $R_{\text{col}}(D, D_{\text{im}}) dD_{\text{im}} n(D) dD$ , where the rate of collisions on to individual planetesimals,  $R_{\text{col}}$ , was defined in equation (11). These collision rates are most strongly affected by the size distribution of the disc and the total amount of cross-sectional area in it, but are largely unaffected by the eccentricities and inclinations of disc material, as  $R_{\text{col}} \propto \sqrt{1 + 1.25e^2/I^2}$ . These orbital parameters do, however, have a strong effect on the outcome of these collisions.

### 6.2 Collisional clump flux

The size distribution of collisional fragments was not discussed in BA99, as it was not permitted by the resolution of their models. Here we use the results of Campo Bagatin & Petit (2001), (hereafter CP01) which considers the size distribution of collisional fragments expected from geometrical constraints. In the CP01 model, a planetesimal of mass  $M$  is broken up sequentially into smaller fragments, where the size of each fragment is determined by the shape of the volume remaining after the previously created fragments have been removed. The physical basis of this approach is that fragmentation proceeds by the coalescence of flaws that propagate through the target after impact. The size distribution they find depends on the mass of the largest remnant created in the collision,  $f_{\text{lr}} = M_{\text{lr}}/M$ . However, apart from the few largest fragments, the distribution was found to follow a power law with an index  $q_c \approx 1.91\text{--}1.97$ , in agreement with their analytical model which predicted a distribution with  $q_c = 1.93$ . This size distribution also concurs with experiments

which find that while  $q_c$  can fall anywhere in the range  $1.6\text{--}2.6$  (Fujiwara et al. 1989; Giblin et al. 1998; Housen & Holsapple 1999), values close to  $1.9\text{--}2$  are more common. While it is not clear why geometrical arguments should apply in the gravity-dominated regime, as here each fragment is a gravitationally reaccumulated conglomerate of the smaller fragments created in the collision, CP01 proved that their model reproduces the size distributions of the asteroid families, the members of which should be reaccumulated fragments (Campo Bagatin et al. 2001), and which have observed slopes of  $q_c = 1.83\text{--}2.17$ .

Here we assume that the size distribution of fragments smaller than the second largest remnant, that with diameter  $D_2$ , follows a power-law size distribution with  $q_c = 1.93$  down to a lower limit  $D_{\text{min,cl}}$ . It may be inappropriate to extrapolate the CP01 size distribution to small fragments, as their distribution may be dominated by different physics to that of larger fragments. For example, Durda & Flynn (1999) found a knee in the size distribution of dust created by the break-up of composite materials due to flaws at material boundaries. This knee occurs at the size of crystals embedded in the planetesimal, ( $\sim 1$  mm in size in their experiment). A flattening of the size distribution for dust smaller than  $1$  mm was also reported by Fujiwara et al. (1989). In any case, we note that while fragments smaller than the radiation pressure blow-out limit,  $D_{\text{min}}$ , may be created in collisions, they would be removed from the disc within an orbital time-scale by radiation pressure. Thus we set  $D_{\text{min,cl}} > D_{\text{min}}$ .

Conserving the volume of the original planetesimal, and scaling the fragment distribution such that there are two fragments larger than  $D_2$ , we find that

$$D_2/D = \left[ \left( \frac{2 - q_c}{q_c} \right) (1 - f_{\text{lr}}) \right]^{1/3}. \quad (17)$$

The total amount of cross-sectional area created by the break-up of a planetesimal is thus

$$\sigma_{\text{col}}(D, D_{\text{im}}) = \sigma_{\text{min}} C_{q_c} [1 - f_{\text{lr}}(D, D_{\text{im}})]^{q_c - 1} (D/D_{\text{min,cl}})^{3q_c - 3}, \quad (18)$$

where  $C_{q_c} = [(2q_c - 2)/(q_c - 5/3)](2/q_c - 1)^{q_c - 1} = 0.323$ . Equation (18) is only valid when  $D_2 < D_{\text{lr}}$ , i.e. when  $f_{\text{lr}} > 1 - 0.5q_c = 0.035$ . When  $f_{\text{lr}} \leq 0.035$  we assume that there is more than one fragment of size  $D_{\text{lr}}$ , and that the  $q_c = 1.93$  distribution applies for smaller fragments. Thus we get a new expression describing the outcome of such a super-catastrophic collision:

$$\sigma_{\text{col}}(D, D_{\text{im}}) = \sigma_{\text{min}} [(2 - q_c)/(q_c - 5/3)] f_{\text{lr}}(D, D_{\text{im}})^{q_c - 2} \times (D/D_{\text{min,cl}})^{3q_c - 3}. \quad (19)$$

Note that the cross-sectional area given in equations (18) and (19) describe that of the debris produced by one of the planetesimals involved in the impact, that which we call the target, and that in both cases this cross-sectional area is  $\propto D_{\text{min,cl}}^{-0.79}$ .

BA99 found that the size of the largest remnant of a collision, whether rubble pile or intact fragment, is determined by the collisional energy in the following way:  $f_{\text{lr}} = 0.5 - s(Q/Q_D^* - 1)$ , where  $s$  is close to  $0.5$  for all materials and impact velocities, and  $Q_D^*$  is the dispersal threshold at the relevant collisional velocity (i.e. not averaged over all impact parameters as before). This is consistent with the results of experiments of catastrophic impacts in the strength regime, which found that  $f_{\text{lr}} = 0.5(Q_D^*/Q)^{1.24}$  (Fujiwara et al. 1989). It is also consistent with the results of experiments of cratering impacts in the strength regime: e.g. Holsapple (1993) used scaling arguments to show that in the strength regime,

$M_{\text{crater}}/M = 1 - f_{\text{ir}} \propto (Q/v_{\text{col}}^2)(v_{\text{col}}^2/Q_S^*)^{1.5\mu}$ ; for  $\mu = 2/3$  and scaling to  $f_{\text{ir}} = 0.5$  when  $Q = Q_S^*$ , we find that  $f_{\text{ir}} = 1 - 0.5Q/Q_S^*$ . For collisions in which  $f_{\text{ir}} > 0.5$ , we use the BA99 results with  $s = 0.5$  so that

$$f_{\text{ir}}(D, D_{\text{im}}) = 1 - 0.5 \left[ \frac{D_{\text{im}}}{X_{\text{ic}}(D)D} \right]^3 \left[ \frac{v_{\text{col}}(D, D_{\text{im}})}{v_{\text{ic}}(D)} \right]^{2-\gamma(D)}, \quad (20)$$

where we remind the reader that  $Q_D^* \propto v_{\text{col}}^{\gamma(D)}$ . For more energetic collisions, however, we use the experimental results summarized in Fujiwara et al. (1989), as these cover a wider range of collisional outcomes, and BA99 only considered collisions in which  $f_{\text{ir}} \approx 0.5$ . Thus for  $f_{\text{ir}} < 0.5$ , we use

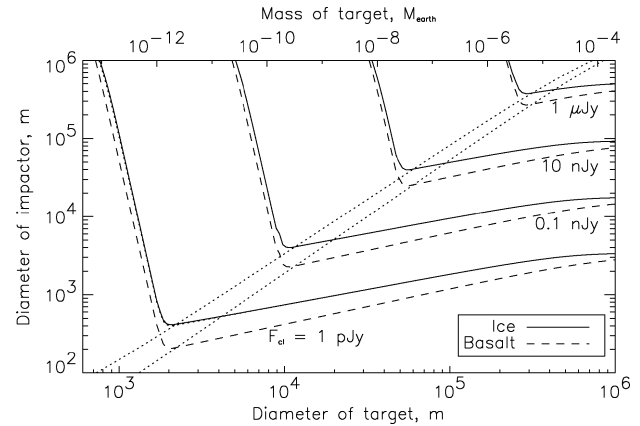
$$f_{\text{ir}}(D, D_{\text{im}}) = 0.5 \left[ \frac{X_{\text{ic}}(D)D}{D_{\text{im}}} \right]^{3.72} \left[ \frac{v_{\text{ic}}(D)}{v_{\text{col}}(D, D_{\text{im}})} \right]^{2.48-1.24\gamma(D)}, \quad (21)$$

which is consistent with the BA99 results in this regime. Here we assume that  $f_{\text{ir}}$  averaged over all impact parameters can be obtained using the averaged  $Q_D^*$  model of Section 5. We also set the physical constraint that  $f_{\text{ir}} \geq (D_{\text{min}}/D)^3$ .

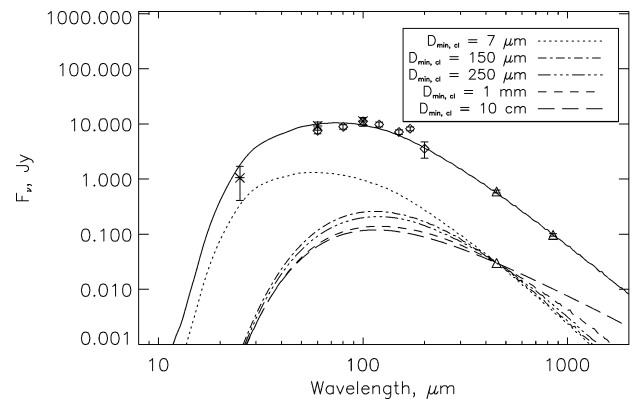
Consider the dust clumps created in collisions between planetesimals in the Fomalhaut disc. In the following discussion, the emission from these clumps is calculated from equation (2) assuming that the composition of this dust is the same as that inferred for the smooth ring from the SED modelling of Section 4, and that this material is spread across the same range of distances from the star as the smooth ring (i.e.  $150 \pm 25$  au). Fig. 9 shows the size of impactor required to produce debris clumps of target material that emit a range of clump fluxes at  $450 \mu\text{m}$ , assuming a fragment cut-off at  $D_{\text{min,cl}} = D_{\text{min}}$ . This shows that for cratering events (i.e. for collisions below the dotted lines), the amount of dust produced depends mostly on the size of the impactor, not on the size of the target; i.e. whether an impactor falls on a planetesimal much larger than itself or one closer to its own size, a similar amount of dust is released from the target. Similarly, the amount of dust resulting from the catastrophic destruction of a target planetesimal (in a collision above the dotted lines) is largely independent of the size of planetesimal impacting the target. It is also interesting to note that debris from the smaller planetesimal in a cratering collision can be brighter than that from the larger planetesimal.

If we had chosen the size distribution to be cut off at  $D_{\text{min,cl}} = 1$  mm, instead of  $7 \mu\text{m}$ , then while the cross-sectional area resulting from a given collision would be decreased by a factor of  $(1000/7)^{-0.79}$ , the flux per unit cross-sectional area would have increased, because while the extra  $7\text{-}\mu\text{m}$  to  $1\text{-mm}$  particles contain a lot of cross-sectional area, they don't emit efficiently at  $450 \mu\text{m}$ . As the  $450\text{-}\mu\text{m}$  flux of the clump in Jy is  $6.5 \times 10^{-3}$  times its cross-sectional area in  $\text{au}^2$  if the cut-off is at  $7 \mu\text{m}$ , but is  $0.11$  times this area if the cut-off is at  $1$  mm, this means that with this larger cut-off, the fluxes in Fig. 9 should be reduced by a factor of  $\sim 3$ .

To determine what the clumps in Fig. 9, which are defined by their  $450 \mu\text{m}$  emission, would emit at other wavelengths, consider Fig. 10, which shows the SED of their emission normalized to  $30 \text{ mJy}$  at  $450 \mu\text{m}$  (the flux of the observed clump). This SED is shown for different  $D_{\text{min,cl}}$  as its shape is strongly dependent on this cut-off due to the increased temperature of small grains (Fig. 4). For example, collisions corresponding to  $1 \mu\text{Jy}$  at  $450 \mu\text{m}$  for a clump with a cut-off at  $7 \mu\text{m}$  (i.e. those close to the top right of Fig. 9) would result in emission of  $11 \mu\text{Jy}$  at  $25 \mu\text{m}$  and  $43 \mu\text{Jy}$  at  $60 \mu\text{m}$ , whereas if fragments from the same collisions had a cut-off at  $D_{\text{min,cl}} = 1$  mm, this would result in emission of  $0.3 \mu\text{Jy}$  at



**Figure 9.** The size of impactor planetesimals in the Fomalhaut disc required to break up the target into fragments that emit a given flux,  $F_{\text{cl}}$ , at  $450 \mu\text{m}$ . The fragments are assumed to follow a power law with  $q_c = 1.93$  down to the radiation pressure blow-out limit of  $7 \mu\text{m}$ . The solid and dashed lines correspond to the models for the collisional properties of Ice and Basalt described in the text; in this target size range Weak Ice has the same properties as Ice. The four lines for each model correspond to  $F_{\text{cl}} = 1 \mu\text{Jy}$ ,  $10 \text{ nJy}$ ,  $0.1 \text{ nJy}$  and  $1 \text{ pJy}$ . The dotted lines mark the limits of threshold catastrophic collisions for the Ice and Basalt models. Collisions below these lines result in cratering from the target planetesimal, while those above result in the complete disintegration of the target.



**Figure 10.** The SED of emission from a clump in Fomalhaut's disc, assuming it is composed of grains with a size distribution with  $q_c = 1.93$  that is cut-off at the small size end below  $D_{\text{min,cl}}$ . The SED of the clump is normalized to  $30 \text{ mJy}$  at  $450 \mu\text{m}$ , which is the flux of the clump observed by Holland et al. 2002, (see Fig. 2c). Also shown is both the observed SED (symbols) and the fit to this SED assuming all this emission comes from the smooth ring (solid line).

$450 \mu\text{m}$ , just  $7 \text{ nJy}$  at  $25 \mu\text{m}$ , and  $0.7 \mu\text{Jy}$  at  $60 \mu\text{m}$ . Note that if the collisional size distribution extends down to the blow-out limit, the clump would appear brighter relative to the smooth ring at mid- to far-IR wavelengths than in the submm.

### 6.3 Clump growth rates

The rate at which the physical extent of a clump grows is determined by the ejection velocity of collisional fragments. This is usually discussed in terms of the parameter  $f_{\text{KE}}$ , which is the fraction of the impact energy,  $E_{\text{col}} = 0.5M M_{\text{im}} v_{\text{col}}^2 / (M + M_{\text{im}})$ , that ends up as kinetic energy of the fragments, a parameter that is not well known. Laboratory experiments of impacts of cm-sized objects imply that the kinetic energy imparted to the largest fragments produced in the collision (i.e. those containing 70–80 per cent of the target mass) is

a fraction  $f_{\text{KE}} = 0.3\text{--}3$  per cent of the impact energy (Fujiwara & Tsukamoto 1980), with an unknown quantity imparted to smaller fragments. Studies of the asteroid families, on the other hand, imply that family-forming events had  $f_{\text{KE}} \approx 0.1$  (e.g. Davis et al. 1989). However, SPH simulations of impacts in the gravity regime imply even higher values of  $f_{\text{KE}} = 0.2\text{--}0.4$  (Love & Ahrens 1996). It has also been reported that  $f_{\text{KE}}$  increases as the velocity of the impact increases  $\propto Q^{0.5}$  (e.g. Fujiwara & Tsukamoto 1980; Paolicchi, Verlicchi & Cellino 1996), and it is expected that collisions with rubble-pile planetesimals have a lower  $f_{\text{KE}}$  than those with solid bodies as a result of the inefficient coupling of impact energy (e.g. Campo Bagatin et al. 2001). While clearly a simplification, in this paper we assume that  $f_{\text{KE}} = 0.1$ , and that this is true in both the strength and gravity regimes. We also assume that this energy is split evenly between the target and the impactor, as this was found to be the case both for cratering events in the strength regime (Hartmann 1988) and for collisions in the gravity regime (Love & Ahrens 1996), although we recognize that the division of this energy depends on the impact parameter (Love & Ahrens 1996), among other things.

The kinetic energy,  $f_{\text{KE}}E_{\text{col}}$ , is not evenly distributed across the fragments. Fragments are found to attain a range of velocities, the distribution of which can be approximated by the relation  $f(v) \propto v^{-k}$ , where  $f(v) dv$  is the fraction of material with velocities between  $v$  and  $v + dv$ . The power-law index,  $k$ , has been reported with values between 3.25 (Gault, Shoemaker & Moore 1963) and 1.5 (Love & Ahrens 1996), where values of  $k < 2$  imply that most of the kinetic energy is carried away by a small mass of material which is moving very fast (e.g. Love & Ahrens 1996, BA99) and which usually originates near the impact site. There also appears to be a dependence of the velocity of a fragment on its mass, in that, on average, lower mass fragments move faster than higher mass ones. The mean velocity of fragments of size  $D$  deduced from laboratory experiments is reported to be  $\propto D^{-r}$ , where  $r$  lies anywhere between 0 (e.g. Giblin et al. 1998) and 0.5 (e.g. Nakamura & Fujiwara 1991), although the scatter about this mean tends to dominate over any strong trend (Giblin et al. 1998). Very little of the kinetic energy is found to be imparted to the largest remnant (Nakamura & Fujiwara 1991; Michel et al. 2001).

Here we calculate the characteristic ejection velocity of target fragments, assuming that the available kinetic energy is distributed among all fragments except the largest remnant; i.e.  $0.5 f_{\text{KE}} E_{\text{col}} = 0.5(1 - f_{\text{r}}) M v_{\text{ej}}^2$ , so that

$$v_{\text{ej}}^2(D, D_{\text{im}}) = \frac{f_{\text{KE}} Q(D, D_{\text{im}}) / [1 + (D_{\text{im}}/D)^3]}{[1 - f_{\text{r}}(D, D_{\text{im}})]}. \quad (22)$$

This means that target fragments created in a cratering event, wherein  $f_{\text{r}}$  is given by equation (20) and  $D_{\text{im}} \ll D$ , have an ejection velocity of  $v_{\text{ej}} \approx \sqrt{2 f_{\text{KE}} Q_{\text{D}}^*}$ . Thus cratering debris from a given target has the same ejection velocity regardless of the size of the impactor. This is because while the amount of kinetic energy imparted to fragments is larger for larger impactors, there is also a greater mass of material ejected from the largest remnant for this to be shared between. This relation also means that the ejection velocity of cratering debris is lowest for that originating from planetesimals 10–1000 m in diameter, as these are the weakest planetesimals (Fig. 6a). Equation (22) also means that the ejection velocity of the fragments of the smaller planetesimal in the collision, for which  $f_{\text{r}} \approx 0$ , is  $v_{\text{ej}} \approx \sqrt{0.5 f_{\text{KE}} v_{\text{col}}(D, D_{\text{im}})}$ ; i.e. the impactor retains a fixed fraction (22 per cent in this case) of its original velocity. This means that for planetesimals for which gravitational focusing is not important, all impactor debris has the same ejection velocity, which is  $\sim 90 \text{ m s}^{-1}$  in the Fomalhaut disc.

In the gravity regime, some of the kinetic energy imparted to the fragments is converted into gravitational energy, as the fragments have to overcome the gravity of the largest remnant. The energy required to disperse all fragments except the largest remnant is  $E_{\text{grav}} = 1.2(GM^2/D)[1 - f_{\text{r}}^{5/3}]$ . Thus the characteristic velocity of the fragments once far from the largest remnant,  $v_{\infty}$ , is given by

$$v_{\infty} = \sqrt{v_{\text{ej}}^2 - v_{\text{grav}}^2}, \quad (23)$$

$$v_{\text{grav}}^2 = 0.4\pi G \rho D^2 [1 - f_{\text{r}}(D, D_{\text{im}})^{5/3}] [1 - f_{\text{r}}(D, D_{\text{im}})]^{-1}, \quad (24)$$

where this assumes that the kinetic energy given to the individual fragments of the largest remnant is small. However, using this model we find that for collisions between the largest planetesimals ( $D > 100 \text{ km}$ ) in the gravity scaling regime in Fomalhaut's disc,  $v_{\text{ej}}$  is sometimes slightly smaller than  $v_{\text{grav}}$ . This occurs when  $Q_{\text{D}}^* < (1/3)\pi\rho G D^2 / f_{\text{KE}}$ . This discrepancy is due to our extrapolation of  $Q_{\text{D}}^*$  from the BA99 models to larger planetesimals and different impact velocities. We already know that the collisional model may not be applicable to the break-up of planetesimals larger than  $\sim 700 \text{ km}$  diameter in this disc, and such collisions are discussed in Section 7. However, it may also be that  $f_{\text{KE}}$  is higher in the gravity regime. Setting  $f_{\text{KE}} = 0.34$  would resolve the discrepancy in this size range. A higher  $f_{\text{KE}}$  may be expected because a much smaller fraction of the impact energy is used in shattering the target; e.g. if  $f_{\text{KE}} \propto (Q/Q_{\text{D}}^*)^{0.5}$  (Fujiwara & Tsukamoto 1980), then threshold catastrophic collisions in the gravity regime would have a very high  $f_{\text{KE}}$  (although some of this kinetic energy would go into the individual fragments of the largest remnant before they reaccumulate, at which point it would be converted into internal energy). Here we account for this discrepancy by artificially increasing  $f_{\text{KE}}$  where required, so that  $v_{\text{ej}} \geq 1.01 v_{\text{grav}}$ . The reasoning behind this is that in the gravity regime the fragments attain a range of velocities, and while those with  $v_{\text{ej}} < v_{\text{grav}}$  remain bound to the largest remnant, most of those that escape must have  $v_{\text{ej}}$  just larger than  $v_{\text{grav}}$ .

The rate at which a clump appears to grow depends on the rate at which the fragments precess around the orbit of the parent planetesimal. The radial and vertical structure of the clump change with time as well, however, the spatial extent of these variations are at or below the level of the azimuthal growth. The precession rate of a fragment is determined by the velocity with which it is ejected in the direction of the orbital motion,  $dv_{\text{k}}$ , as this causes a change in the semimajor axis of the fragment's orbit, and hence in its orbital period,  $dt_{\text{per}}/t_{\text{per}} = -3 dv_{\text{k}}/v_{\text{k}}$ . The ejection velocity field is generally found to be isotropic, both in experiments (Giblin et al. 1998) and analysis of the orbital elements of the asteroid families (Zappala et al. 1996). Thus the average velocity of all fragments in the direction of orbital motion is  $dv_{\text{k}} \approx v_{\infty}/\sqrt{3}$ , and the rate of growth of the total azimuthal extent of the clump is

$$\dot{\theta} = 2\sqrt{3}v_{\infty}/a. \quad (25)$$

Before continuing, we define the characteristic lifetime of a clump of target fragments created in a collision between the target (of size  $D$ ) and an impactor (of size  $D_{\text{im}}$ ),  $t_{\text{cl}}(D, D_{\text{im}})$ , to be equal to the average time it takes for the clump fragments to occupy half the ring (i.e.  $\pi/\dot{\theta}$ ):

$$t_{\text{cl}}(D, D_{\text{im}})/t_{\text{per}} = (1/4\sqrt{3})[v_{\text{k}}/v_{\infty}(D, D_{\text{im}})]. \quad (26)$$

Thus in this model impactor debris from collisions between planetesimals in Fomalhaut's disc remains in a clump for about six orbital periods, while target debris clumps last for between 10 and 700

orbital periods, depending on the dispersal threshold of the target (clumps from weak, 10–1000 m planetesimals last longest).

#### 6.4 Discussion

Combining the collision rate from Section 6.1 and the clump growth rate from Section 6.3, we find that the number of clumps we expect to see at any one time that are smaller in azimuthal extent than  $\theta$ , and that were created in collisions between (target) planetesimals of size  $D$  to  $D + dD$  and those (impactors) of size between  $D_{\text{im}}$  and  $D_{\text{im}} + dD_{\text{im}}$  is given by

$$N(<\theta) = R_{\text{col}}(D, D_{\text{im}}) dD_{\text{im}} n(D) dD(\theta/\dot{\theta}). \quad (27)$$

We also know how bright those clumps should be, as this was defined in Section 6.2 (Fig. 9). Thus we can integrate over all impacts that create clumps brighter than a specific limit to get the total number of clumps we expect to see in the disc that are smaller in azimuthal extent than  $\theta$  and brighter than  $F_{\text{cl}}$ ,  $N(<\theta, >F_{\text{cl}})$ .

Clumps that are greater than  $\sim 180^\circ$  in extent are no longer clumps, but radially confined rings that blend into the background disc. Furthermore, the emission from such rings could be affected by the subsequent collisional evolution of their constituents.<sup>4</sup> Here we define the *clump function* to be the number of clumps we expect to see in the disc that are smaller than  $180^\circ$  in extent and brighter than  $F_{\text{cl}}$ ,  $N(<180^\circ, >F_{\text{cl}})$ . The number of clumps brighter than this limit, but with a maximum extent smaller than  $180^\circ$ , can then be determined using the relation:

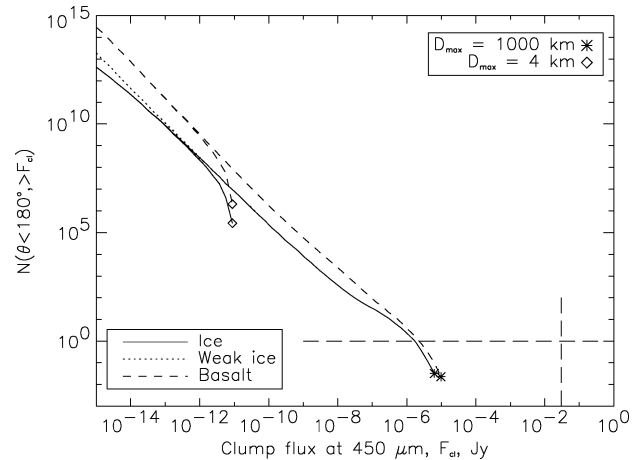
$$N(<\theta, >F_{\text{cl}}) = N(<180^\circ, >F_{\text{cl}}) \times (\theta/180^\circ). \quad (28)$$

In Fig. 11 we have plotted the clump function in Fomalhaut's disc assuming that the disc comprises: (i) just the planetesimals in the collisional cascade; (ii) the collisional cascade plus an additional  $\sim 1M_{\text{J}}$  component of primordial planetesimals with a size distribution defined by  $q_{\text{p}} = 1.84$  extending up to those 1000 km in diameter. The steepness of the clump function in this plot, combined with the linear growth rate of clumps, implies that this clump function can be loosely interpreted as the number of clumps we expect to see of flux  $\sim F_{\text{cl}}$  that are  $90^\circ$ – $180^\circ$  in extent. The origin of these clumps could be in any collision on the appropriate line in Fig. 9. However, we find that clumps produced in the gravity regime are most likely to be target debris<sup>5</sup> from a threshold catastrophic collision; e.g. any  $1\text{-}\mu\text{Jy}$  clump we observe is most likely material created in a collision between two bodies  $\sim 300$  km in diameter.

We define the *intrinsic clumpiness* of a disc as the flux at which the clump function,  $N(<180^\circ, >F_{\text{cl}}) = 1$ . This can be interpreted as the approximate flux of the largest single clump we would expect to see in the disc. However, as a disc contains a distribution of clump sizes, smaller fainter clumps would also be observed down to a level determined by the beam size. Observations with a beam of full width at half-maximum FWHM (in arcsec), would see all clumps with  $\theta < \theta_{\text{FWHM}}$ , where  $\theta_{\text{FWHM}} = 360^\circ \times (\text{FWHM}r_{\star}) / (2\pi r_{\text{cl}})$  ( $r_{\text{cl}}$  is the distance of the clump from the star in au) as unresolved point sources; those with  $\theta_{\text{FWHM}} < \theta < 180^\circ$  would be resolved, while

<sup>4</sup>The subsequent collisional evolution of collisional fragments would not affect their emission by much until  $\theta > 180^\circ$ , as the collisional lifetime of this material, assuming it is the same as that of the background disc (Fig. 7), is much longer than the age of the clump (equation 26). This evolution would tend to reduce the brightness of the clump, as collisions remove the smallest dust first.

<sup>5</sup>As target and impactor debris remains in a clump for different lengths of time, we have treated the two components separately.



**Figure 11.** The clump function,  $N(\theta < 180^\circ, >F_{\text{cl}})$ , of the Fomalhaut disc, i.e. the number of clumps we expect to see that are brighter than  $F_{\text{cl}}$  but smaller than  $180^\circ$  in azimuthal extent. The solid, dotted and dashed lines correspond to the models for the collisional properties of Ice, Weak Ice and Basalt described in the text. The two lines for each model assume that the size distribution with  $q_{\text{d}} = 1.84$  extends up to maximum planetesimal sizes of 4 km and 1000 km. The intrinsic clumpiness of the disc [the  $F_{\text{cl}}$  at which  $N(\theta < 180^\circ, >F_{\text{cl}}) = 1$ ] is shown by the horizontal long-dashed line. Its observed clumpiness,  $F_{\text{cl}} = 30$  mJy, is shown by the vertical long-dashed line. For collisions to be the likely cause of this clumpiness we would expect the clump function to pass through the intersection of these lines.

larger ones would be rings rather than clumps. Thus, the number of clumps we would expect to see as point sources above a certain flux would be  $\theta_{\text{FWHM}}/180^\circ$  times the clump function, and the brightest unresolved point source we could expect to see would be the flux at which the clump function,  $N(<180^\circ, >F_{\text{cl}}) = 180^\circ/\theta_{\text{FWHM}}$ .

Fig. 11 shows that if Fomalhaut's disc is comprised solely of planetesimals in the collisional cascade, its largest (4 km) planetesimals are so numerous, and their clumps last so long, that at any one time we would expect to see roughly  $10^6$  clumps resulting from their destruction that are  $<180^\circ$  in extent. Thus even if the sensitivity of our observations was sufficient to detect these clumps (which would emit just  $\sim 10$  pJy), the disc would appear smooth unless our beam size was also  $<60$   $\mu\text{as}$ ! However, if the disc also contains a primordial planetesimal distribution as described in (ii) above, the intrinsic clumpiness of the disc would be  $\sim 2$   $\mu\text{Jy}$  (Fig. 11). Brighter clumps could also exist in the disc (up to 10  $\mu\text{Jy}$ ), however, such large collisional events happen infrequently. For these clumps, the clump function defines the probability of our witnessing a clump that bright and smaller than  $180^\circ$ ; i.e. at any one time we have a 1:100 chance of seeing a 10- $\mu\text{Jy}$  clump (if the proposed distribution is correct).

Could intrinsic clumpiness be the likely cause of the clump observed in Fomalhaut's disc (Section 2)? If so, we would expect the derived intrinsic clumpiness at 450  $\mu\text{m}$  to be 30 mJy; i.e. the clump function should pass through  $N(<180^\circ, >30 \text{ mJy}) \approx 1$ . This is not true in either of the examples shown in Fig. 11. However, we have no reason to expect that  $q_{\text{p}} = 1.84$ , nor that this distribution extends up to  $D_{\text{max}} = 1000$  km. A lower  $q_{\text{p}}$  would increase the number of 1000-km planetesimals in the disc, and so increase the number of clumps expected from their destruction (although this would also imply a disc more massive than  $1M_{\text{J}}$ , see equation 16). The clump function could also be underestimated due to uncertainties in our model. Particularly, the lifetime of clumps produced by the destruction of planetesimals in the gravity regime is not well known (as

discussed in Section 6.3), and may have been underestimated. These factors could mean that the intrinsic clumpiness is actually at the level of the break-up of the largest planetesimal in the disc. However, if  $D_{\max} = 1000$  km, this is just  $\sim 10 \mu\text{Jy}$ , some 3000 times lower than that required for intrinsic clumpiness to explain the observed clump.

However, it is possible that our model also underestimates the amount of dust produced in collisions in the gravity regime, as smaller dust grains might be predominant in the escaping fragments from these collisions. If the velocity imparted to the fragments is size-dependent, and smaller fragments attain higher velocities, then it would be the largest fragments of the original distribution that reaccumulate, while the smaller ones are more likely to be released. Furthermore, the surfaces of large planetesimals are expected to be covered with large quantities of small dust grains in the form of a regolith (see Section 7.2); these grains would be preferentially released in a collision. It has not been possible to include such effects in our model because the outcome for  $\mu\text{m}$ -sized dust from km-scale impacts has yet to be studied. Based on arguments outlined in Section 7.1, the most our model could be underestimating the flux by is a factor of 1000.<sup>6</sup> However, the real value is likely to be much less than this. In any case, collisions among non-gravitationally focusing planetesimals result in insufficient flux for them to be the cause of the observed clump.

### 6.5 For future consideration

Could Fomalhaut's intrinsic clumpiness be observed with future instrumentation? With a beam size of 0.6 arcsec, Fig. 11 implies that we could expect to observe clumps in Fomalhaut's disc between  $\sim 0.1 \mu\text{Jy}$  (unresolved point sources) and  $2 \mu\text{Jy}$  (resolved clumps) at  $450 \mu\text{m}$ . Depending on the amount of  $\mu\text{m}$ -sized dust produced in collisions, these limits correspond to clumps of between  $1 \mu\text{Jy}$  and  $0.02 \text{ mJy}$  at  $25 \mu\text{m}$  (Fig. 10). A 0.6-arcsec beam is close to that of the *ALMA* in compact configuration at  $450 \mu\text{m}$ , or that of a mid-IR instrument operating at  $25 \mu\text{m}$  on an 8-m telescope. While the clumpiness of Fomalhaut's disc proposed above is at least an order of magnitude below the anticipated sensitivities of such instruments, the above figures include uncertainties (such as the potential factor of 1000 underestimate of the amount of dust created in a collision, Section 6.4) which could render these clumps detectable. If so, their observed magnitude and size distribution could be used to constrain the primordial planetesimal population.

Also, we note that this type of clumpiness could provide confusion for a mission to find terrestrial planets using space-based  $10\text{-}\mu\text{m}$  interferometry, such as the designs proposed for *DARWIN* or *TPF* (e.g. Mennesson & Mariotti 1997). The biggest potential impact of disc clumpiness on such a mission would come from a warm zodiacal cloud-like disc  $\sim 3$  au from the star; this topic is discussed in Wyatt (2001). However, here we point out that if viewed at the distance of Fomalhaut, the Earth would have a  $25\text{-}\mu\text{m}$  flux of  $\sim 0.8 \mu\text{Jy}$ , while the primordial planetesimal distribution proposed in Fig. 11 implies that the Fomalhaut disc could contain a comparable  $\sim 0.2\text{-}\mu\text{Jy}$  clump that would be unresolved in a  $0.05\text{-arcsec}$  beam, which would be that of a 100-m baseline interferometer at this wavelength.

<sup>6</sup>The collisional model predicts that the destruction of a 1000-km planetesimal would result in dust that emits a flux of  $10 \mu\text{Jy}$ , whereas the maximum possible emission from a pulverised 1000-km body, i.e. when it is ground into  $150 \mu\text{m}$ – $1 \text{ mm}$  dust, is  $10 \text{ mJy}$  (Fig. 12b).

## 7 THE COLLISIONAL CLUMP HYPOTHESIS

Consider the magnitude of collision required to produce the clump observed in the Fomalhaut disc. Section 6.4 showed that this clump cannot have been produced by the intrinsic clumpiness of the collisional cascade, or by the break-up of the primordial planetesimal population, because such collisions do not produce enough dust. Thus we start by determining the smallest planetesimal that could have produced the clump using a simple model in which the clump is comprised of grains of just one size.

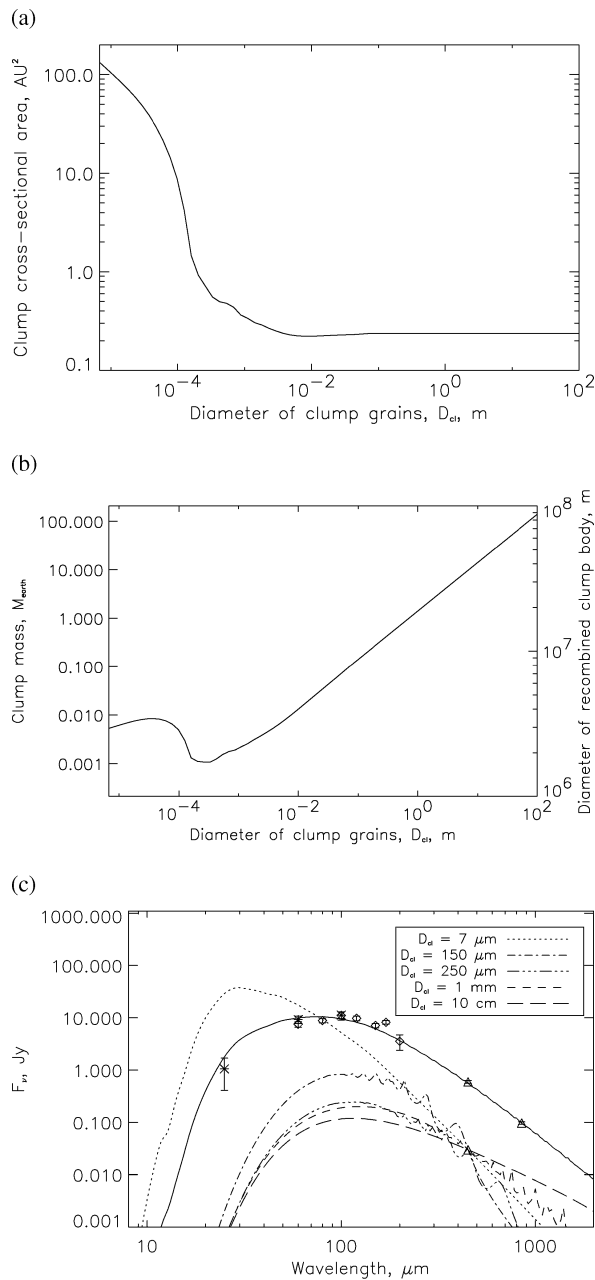
### 7.1 Single size dust grained clump

If the dust grains in the clump are all of the same size,  $D_{\text{cl}}$ , then Fig. 12(a) shows how the observed  $450\text{-}\mu\text{m}$  flux of the clump,  $F_{\nu}(450 \mu\text{m}) = 30 \text{ mJy}$ , determines the total cross-sectional area of material in the clump,  $\sigma_{\text{cl}}$  (found by inverting equation 2), for different  $D_{\text{cl}}$ . Thus the projected area of this clump must be at least  $0.2 \text{ au}^2$ , much larger if a significant fraction of the material is in dust smaller than  $1 \text{ mm}$ . Fig. 12(b) also shows how this flux determines the total mass of material in the clump, where  $m_{\text{cl}} = (2/3)\rho\sigma_{\text{cl}}D_{\text{cl}}$ . Thus the clump must have a mass of at least a thousandth of that of the Earth, and would only be this small if all the grains in the clump are between  $150 \mu\text{m}$  and  $1 \text{ mm}$  in diameter. If there is a significant fraction of mass in particles smaller than  $150 \mu\text{m}$  then we would infer a more massive clump, as these particles do not emit efficiently at  $450 \mu\text{m}$ . Similarly, we would infer a more massive clump if there is a significant fraction of mass in particles larger than  $1 \text{ mm}$ . It would be possible to constrain the size of particles in the clump if the clump could be imaged at shorter wavelengths (see Fig. 12c). However, for now we can rule out the possibility that the clump consists of grains that are all smaller than  $\sim 100 \mu\text{m}$ , as the emission from such a clump would exceed that observed, presumably from the smooth ring, at mid- to far-IR wavelengths.

Thus if the material in this clump originated from a single body, that planetesimal would have to have been at least  $1400 \text{ km}$  in diameter. To understand how such a large body, which must have formed through runaway growth, and for which gravitational focusing is important, interacts with the planetesimal disc, we start by considering how such a planetesimal grew in the first place.

### 7.2 Runaway planetesimal growth

The runaway population are the largest members of the distribution of planetesimals growing in the protostellar nebula. They grew initially in the same way as the rest of the planetesimals, by the accretion of nearby, usually much smaller planetesimals. However, kinetic energy exchange in encounters with other planetesimals (i.e. dynamical friction) led to larger planetesimals having more circular orbits (Wetherill & Stewart 1993). The consequent low relative velocity and high gravitational focusing factor for collisions with large bodies (e.g. equation 13), meant that large bodies (once larger than  $\sim 1 \text{ km}$ ) grew more rapidly than smaller planetesimals, creating a bimodal size distribution – an accumulation cascade and a population of large, so-called runaway, planetesimals. This runaway growth would have been relatively short, typically several Myr (e.g. Kenyon 2002), and would have been halted once the runaways became large enough to perturb the disc just enough to affect their collisional velocities (i.e. once larger than a few  $100 \text{ km}$ ). The runaways then continue to grow, but at a slower rate than smaller planetesimals; so-called oligarchic growth (e.g. Kokubo & Ida 2000). Once these runaways are large enough, however, their gravitational



**Figure 12.** The total cross-sectional area (a) and mass (b) of material in Fomalhaut’s clump, as well as the SED of this clump (c), assuming it is comprised of grains of just one size,  $D_{cl}$ , and setting the clump’s emission to 30 mJy at 450  $\mu\text{m}$ . As in Section 6, the clump is assumed to be composed of the same material as that inferred for the smooth ring and the clump’s emission is calculated assuming that this material is spread across the same range of distances from the star as the smooth ring. Also shown in (b) is the diameter of the body that the clump mass would correspond to when recombined into a single body, and (c) also shows the observed SED (symbols) and the fit to this SED assuming all this emission comes from the smooth ring (solid line).

perturbations increase the impact velocities of members of the accumulation cascade so that collisions result in their net destruction, rather than in their net accretion.

During this oligarchic growth, the runaways try to remain dynamically isolated from each other. However, this is only possible as long as their separation is at least 5 Hill’s radii (Kokubo & Ida 1995), where the Hill’s radius of a planetesimal is  $r_H = a(M_{pl}/3M_\star)^{1/3}$ . The

increasing mass of a runaway means that any isolation is not permanent and interactions with other runaways are inevitable. These interactions would result in both collisions amongst the runaway population, as well as increases in the eccentricities of this population, thus increasing their feeding zones (e.g. Weidenschilling et al. 1997) and making mutual collisions more likely. Their growth would eventually be halted once they had scattered or accreted all of the planetesimals in their feeding zone. The angular momentum exchange caused by this scattering could cause the orbit of the newly formed planet to migrate either towards or away from the star (e.g. Hahn & Malhotra 1999).

A runaway body would have started as an incoherent body built up of layers of regolith. Mutual collisions between runaways would have resulted in the runaway often being completely fragmented, then reaccumulating. Because smaller fragments attain higher ejection velocities, the reaccumulated body would be sorted according to grain size with the larger grains towards the centre and fine grains on the surface (e.g. Britt & Consolmagno 2001). Impacts of smaller bodies on to the runaway would result in the comminution of the surface regolith layer of the runaway (e.g. McKay, Swindle & Greenberg 1989). Planetesimals in the solar system have such regoliths: the Moon’s regolith is 5–10 m deep in the mare region, and mostly comprises (80–90 per cent) dust grains smaller than 1 mm in size (e.g. Heiken 1975), while that of 20-km diameter Phobos is thought to be more than 100 m thick in places (Veverka & Thomas 1979), and the regolith layer on 200-km sized asteroids is predicted to extend to a few km in depth (Housen & Wilkening 1982). Thus we expect runaways to be composed of a megaregolith structure on top of which is a deep dusty regolith layer.

### 7.3 Observing the final throes of planet formation?

Within this scenario there are three ways in which the observed dust clump could be associated with a runaway planetesimal in the ring: (i) it is unbound material generated in a recent collision between two runaways; (ii) it is bound material constantly replenished by impacts into a growing runaway; (iii) it is an infalling envelope of dust that is being accreted from the collisional cascade on to the runaway. The clump could also be associated with a planet orbiting interior to the ring: (iv) its dust could have been created by the break-up of planetesimals that are trapped in resonance with the planet.

#### 7.3.1 Unbound runaway–runaway collisional clump

If we assume that the  $M_\oplus/1000$  of dust that we see originated in just one collision, then it must have originated from a parent planetesimal more than 1400 km in diameter. For this dust to be released from the gravity of the parent body, the impacting planetesimal must have had enough mass that the kinetic energy imparted to the fragments exceeded the escape velocity. Because impacts with gravitationally focusing planetesimals take place at the escape velocity, at best an impactor can only release  $f_{KE}$  times its own mass. Thus the observed clump, if the product of one collision, must have involved the collision of two runaway planetesimals at least 1400 km in diameter. In fact, the planetesimals may have to have been significantly larger than this, as experiments show that transmission of kinetic energy is much less efficient for collisions into sand than into rock (e.g. Holsapple 1993) leading some authors to speculate that  $f_{KE} \approx 0.01$  for collisions with such rubble piles (Campo Bagatin et al. 2001).

How likely are we to be observing a collision between two runaways? Planet formation simulations do predict the formation of a

significant population of runaways. Also, we know that runaway–runaway collisions do occur: based on the similarity of the abundances of isotopes of oxygen in lunar samples to those on the Earth, Wiechert et al. (2001) concluded that the Mars-sized impactor that collided with the Earth to form the Moon formed at the same heliocentric distance as the Earth (i.e. it was a nearby runaway). To have grown to  $>1400$  km, the runaways in Fomalhaut’s disc must have grown in isolation from each other, with at least  $5r_H$  separation. This means that there can be at most  $100 M_{\oplus}/1000$  runaways in Fomalhaut’s 50-au wide ring. As each collision would result in the merger of the runaways and so in the loss of one of the runaways, this leaves at most about 100 possible collisions resulting in the observed clump, that is if the runaways are not scattered out of the disc by their mutual perturbations before they collide.

The length of time this material would remain in a clump depends on its ejection velocity. The only studies that can throw light on the outcome of such collisions are the SPH simulations of the formation of the Moon (e.g. Cameron 1997; Canup & Asphaug 2001) and of the asteroid families (Michel et al. 2001). However, even these do not predict the outcome for collision fragments that are smaller than their resolution ( $>km$ ), or indeed what would happen to the dusty regoliths on the colliding bodies. What they do show, however, is that in a collision between two runaways a significant mass of fragments does escape their gravitational field (see Table 1 of Canup & Asphaug 2001). Thus such collisions would result in the formation of an unbound clump. Also the ejection velocities of fragments from collisions between massive, but non-gravitationally focusing, planetesimals imply a clump lifetime (equation 26) of 20–30 orbital periods for collisions in the asteroid belt (Michel et al. 2001). Thus if this clump lifetime is also typical for runaway–runaway collisions in Fomalhaut’s disc, we could only expect to see a clump from their mutual collisions for a total of  $\sim 3$  Myr of the life of the disc.

This is an extremely short window to be witnessing the product of runaway–runaway collisions, given the age of the system (3 Myr is 1–2 per cent of the life of the disc). However, the lifetime of a clump created in a collision between gravitationally focusing bodies in Fomalhaut’s disc could be an order of magnitude longer than those of clumps created in the asteroid belt. Also, planet formation models predict that the growth of the runaways to sizes large enough both to cause the observed clump and to ignite of the collisional cascade can take a few hundred Myr at the distance of Fomalhaut’s disc (e.g. Kenyon & Bromley 2001). Thus we leave this topic to be explored using more detailed planet formation and collision models, but remark that it appears very unlikely at this stage that runaway–runaway collisions are the cause of the observed clump.

### 7.3.2 Bound collisionally replenished envelope (CRE)

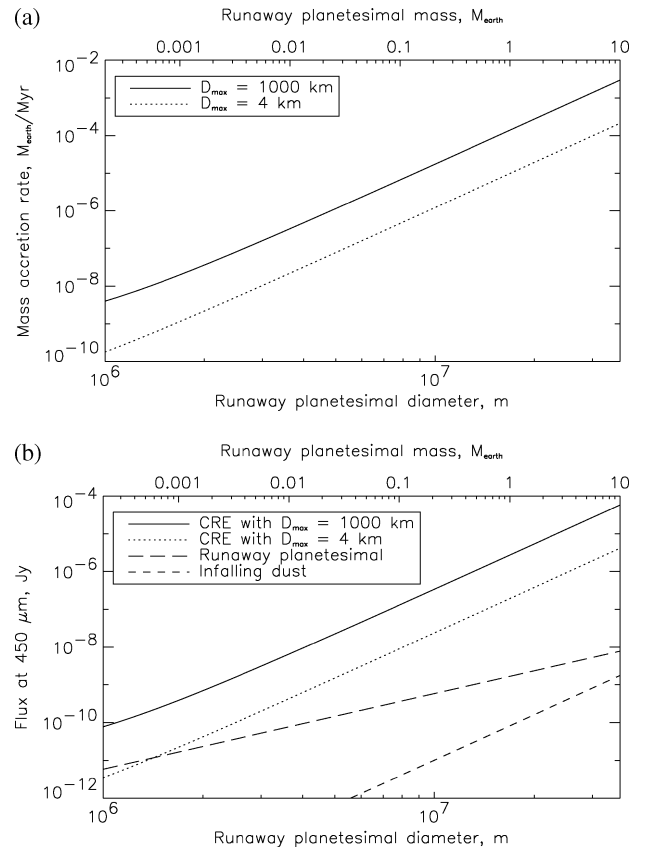
Impacts into a runaway planetesimal would launch a cloud of regolith dust from its surface. The majority of this dust would then collapse back on to the runaway. The maximum distance out to which dust would remain bound to the runaway would be to its Hill’s radius, as outside this its gravity would be weaker than that of the star; this corresponds to  $\sim 1.2$  au for a  $1-M_{\oplus}$  planetesimal orbiting at 150 au around Fomalhaut. Thus for the observed dust clump, which has a projected area of  $>0.2$  au<sup>2</sup> (Section 7.1), to be bound to its parent planetesimal, that planetesimal would have to have a mass of at least  $0.01M_{\oplus}$ . Here we consider whether Fomalhaut’s clump could be caused by repeated collisions on to a runaway from non-

gravitationally focusing planetesimals sustaining a dusty envelope of regolith material extending to the Hill’s radius. Because all collisions with runaways occur at the escape velocity, and the ejection velocity required for dust to reach to the Hill’s radius is close to (just  $5 \times 10^{-5}$  times less than) that escape velocity, energy constraints imply that the maximum amount of regolith dust that an impacting planetesimal could launch to such a distance would be equal to its mass. Thus the maximum mass of material in such an envelope at any one time is equal to the mass accretion rate on to the runaway times the maximum length of time that the bound dust could remain in the envelope.

The mass accretion rate of a runaway of size  $D_{run}$  is given by

$$\dot{M}_{accr}(D_{run}) = \int_{D_{min}}^{D_{max}} (\pi\rho/6) D_{im}^3 R_{col}(D_{run}, D_{im}) dD_{im}. \quad (29)$$

This is plotted in Fig. 13(a) for size distributions that extend up to 4 km (i.e. the collisional cascade) and 1000 km (i.e. including a hypothetical primordial planetesimal population). Note that as the growth rate of runaways in the current disc is relatively slow, they must have formed when the eccentricities and inclinations of disc material were much lower, as then the accretion rate would have been higher ( $\dot{M}_{accr} \propto [If(e, I)]^{-1}$ ). The maximum amount of time dust from any one collision would remain in the envelope would



**Figure 13.** The interaction of a runaway planetesimal with the non-runaway population in Fomalhaut’s disc: (a) the rate of accretion of mass on to the runaway; (b) the flux of clump expected from this interaction due to both the collisionally replenished envelope (CRE) of regolith material and the infalling dust envelope, as well as the emission from the runaway itself (assuming a temperature of  $\sim 45$  K). Because a runaway accretes most of its mass from the largest members of the population, the mass accretion rate and the flux from the CRE are shown for both  $D_{max} = 4$  km and 1000 km (dotted and solid lines respectively).



be twice the free-fall time-scale from the Hill's radius (e.g. Love & Ahrens 1996), i.e. about half an orbital period, or 650 yr at 150 au from Fomalhaut. Some material could, however, remain for longer if injected into orbit around the runaway (e.g. Canup & Asphaug 2001; Michel et al. 2001).

The maximum 450- $\mu\text{m}$  flux a given mass of material in Fomalhaut's disc could emit can be calculated from Fig. 12(b). If all this mass was in dust 150  $\mu\text{m}$  to 1 mm in size, this would emit 30 mJy/0.001 $M_{\oplus}$ . Fig. 13(b) uses this to show the maximum flux of such a collisionally replenished envelope (CRE). This shows that even the CREs surrounding 10- $M_{\oplus}$  planets could not explain the observed dust clump. However, the flux of these CREs could surpass that of the runaway itself by up to a factor of 10 000 for the 450- $\mu\text{m}$  flux of a 10 $M_{\oplus}$  planet (or perhaps even higher at shorter wavelengths, see Fig. 12c). These CREs could also be bright enough that it is the distribution of massive runaways that determines the intrinsic clumpiness of a disc, rather than the distribution of non-gravitationally focusing planetesimals (Section 6.4).

### 7.3.3 Infalling dust envelope

Another reason why there would be an enhanced density of dust in the vicinity of a runaway is that it is accreting dust from the disc's collisional cascade. The total cross-sectional area of material in this infalling envelope is approximately the rate of accretion of cross-sectional area of the runaway, given by

$$\dot{\sigma}_{\text{accr}}(D_{\text{run}}) = \int_{D_{\text{min}}}^{D_{\text{max}}} (\pi/4) D_{\text{im}}^2 R_{\text{col}}(D_{\text{run}}, D_{\text{im}}) dD_{\text{im}}, \quad (30)$$

multiplied by the infall time-scale, which is the freefall time-scale from its Hill's radius (one quarter of an orbital period). The flux from this infalling material can then be calculated by considering that the disc emission of 0.595 Jy at 450  $\mu\text{m}$  (Table 1) is predicted to originate from  $\sigma_{\text{tot}} = 33.7 \text{ au}^2$  of material (Section 4.3). The resulting flux is plotted in Fig. 13(b), which shows that such an infalling envelope can be ignored as it is fainter even than the runaway.

### 7.3.4 Dusty resonant ring

While no planets have been detected around Fomalhaut, the existence of a planetary system orbiting within the observed dust ring would naturally explain the lack of dust emission there, as such a system would have cleared the planetesimal disc within which it formed. As stated in Section 7.2, this clearing could have caused the planets to migrate outwards. In the course of this migration, the outer mean motion resonances of the outermost planet of the system would have swept through the disc of planetesimals outside the orbit of the planet. Many of these planetesimals would have become trapped in these resonances and would then have migrated out with the planet, while those that did not get trapped would have been scattered by the approaching planet. This is the mechanism which is invoked to explain the large number of Plutinos, Kuiper belt objects that are trapped in Neptune's 2:3 resonance (e.g. Jewitt 1999). Hahn & Malhotra (1999) estimated that the required migration of Neptune's orbit from 23 to 30 au could have been caused by the clearing of a 50- $M_{\oplus}$  planetesimal disc over the first 50 Myr of the solar system. We note that these last two figures are comparable to the inferred mass and age of Fomalhaut's disc; thus it could be argued that similar migration and trapping are to be expected in this system had a planet formed

in a (now scattered) planetesimal disc interior to the one we see today.

The paths of eccentric resonant orbits when plotted in a frame co-rotating with the mean motion of the planet exhibit loops which occur at the pericentres of the orbits (see Fig. 8.4c of Murray & Dermott 1999). As only those resonant orbits that have their conjunction with the planet when their apocentres are stable (at least when the planet has a low eccentricity), these loops only occur at specific locations relative to the planet. Planetesimals that are evenly distributed within these stable resonances would be most densely concentrated in these loop regions. Thus, assuming that the fragments created by the collisional break-up of such resonant planetesimals also remain in resonance, the resulting dust ring would be clumpy (Wyatt 1999). The number and distribution of these clumps would be determined by the fraction of planetesimals that are trapped in the different resonances. The one clump in Fomalhaut's ring implies that the material is trapped in a 1:2 resonance, although other clumps could be hidden in the lobes. Very few Kuiper belt objects are observed to be trapped in Neptune's 1:2 resonance, however this may be an observational limit, or because the primordial Kuiper belt did not extend out that far, and/or Neptune's orbit did not migrate far enough. The fraction of planetesimals that we require to be trapped in the 1:2 resonance to cause the observed 30 mJy clump would be of the order of 5 per cent, as resonant material can spend a significant fraction of its time in the clump region (the actual fraction is determined by its eccentricity) and while in the clump this material would be closer to the star and therefore hotter and brighter than that in the rest of the ring.

If this interpretation is correct we would predict that the planet orbits at 80 au with a period of about 500 yr, and that it is currently located on the opposite side of the star from the dust clump. We would also predict that the clump would orbit the star with the planet (direction unknown), causing motion on the sky of up to 0.2 arcsec yr<sup>-1</sup>, and that the clump would be hotter than the rest of the ring. A model for this resonant structure, and the constraints it sets upon the mass of the planet, and that of the original planetesimal disc, is beyond the scope of this paper and will be discussed in a forthcoming paper.

We note that the resonant ring we propose here is formed in a different manner to that proposed to explain the Vega and  $\epsilon$  Eridani dust clumps (Ozernoy et al. 2000; Wilner et al. 2002). The dust in their rings is trapped into planetary resonances through inward dust migration caused by P-R drag, rather than by outward planetary migration. Thus their rings are akin to the Earth's resonant ring (Dermott et al. 1994). We contend that this mechanism would be inefficient in this system, as grains would be destroyed by collisions before they migrate into the resonances by P-R drag (e.g. Section 3.3; WDT99; Wyatt 1999).

## 8 CONCLUSIONS

We have presented a detailed study of collisional processes across the whole size range in material in the Fomalhaut debris disc. This is the first such study of a debris disc that we are aware of. We have shown that the spatial distribution of the dust derived from the 450- $\mu\text{m}$  image of the disc, combined with modelling of its SED, are sufficient to constrain the size distribution of the dust but not to set stringent constraints on its composition. The size distribution we inferred is the same as that expected from a disc in collisional equilibrium, although we recognize that our parametrization of this distribution is only an approximation. The emission that we observe originates from material in the size range from 7  $\mu\text{m}$  to

0.2 m, however collisional lifetime arguments imply that this material originated in a cascade involving the break-up of planetesimals up to 4 km in size. The inferred mass of material in this collisional cascade is 20–30  $M_{\oplus}$ . We speculate that there may be a population of primordial planetesimals with sizes between 4–1000 km, however these collide too infrequently for their collisional debris to have contributed to the dust we observe today. The slope of the size distribution of these planetesimals could be similar to that of the cascade and extend up to 1000 km, based on models of planet formation. Dust clumps caused by the break-up of these planetesimals would be at too low a level to be detected with current technology, however it may be possible to observe these clumps in the future (e.g. with the *ALMA*). With such observations we would be able to constrain the population of these large planetesimals – indeed this may be the only way of detecting their presence. There may also be a population of runaway planetesimals larger than 1000 km which would still be growing. It may be possible to detect these directly, as they would always be enveloped by a cloud of dust launched from their regolith by repeated collisions. These dust envelopes could also be at a level which it may be possible to observe in the future.

We propose two possible origins for the 30 mJy clump that is observed in Fomalhaut’s disc. First, it could have been created in a collision between two runaway planetesimals. We are limited in being able to assess the likelihood of this by uncertainties in the outcome of collisions between such massive bodies, as well as by uncertainties in the number and collision frequencies of such bodies at the end of planet formation. Our best estimate, however, shows that this possibility is unlikely unless both the formation of the runaways and the ignition of the collision cascade occurred within the last few Myr. Another possibility is that  $\sim 5$  per cent of the planetesimals in the ring were trapped in 1:2 resonance with an inner planet when it migrated out as a result of the clearing of a residual planetesimal disc. This hypothesis is strengthened by the fact that the mass and age of Fomalhaut’s disc are comparable to that required for Neptune’s migration and the consequent resonant trapping of Kuiper belt objects. Constraints on the mass of Fomalhaut’s hypothetical planet will be addressed in a future paper. We predict that the orbital motion of such a resonant clump ( $0.2 \text{ arcsec yr}^{-1}$ ) could be detectable within a few yr, and that this clump would be more prominent at shorter wavelengths.

It is clear from this study that collisions are fundamental to the formation and evolution of debris discs, not only as the mechanism which creates the dust we see, but as a cause of potentially observable structure in the disc. Unless future collisional and planet formation models can prove otherwise, it appears unlikely that Fomalhaut’s observed clump is collisional in origin. However, this does not rule out a collisional origin for the clumps seen in other debris discs. These discs will be the subject of a future paper.

## REFERENCES

- Arakawa M., 1999, *Icarus*, 142, 34  
 Augereau J. C., Lagrange A. M., Mouillet D., Papaloizou J. C. B., Grorod P. A., 1999, *A&A*, 348, 557  
 Augereau J. C., Nelson R. P., Lagrange A. M., Papaloizou J. C. B., Mouillet D., 2001, *A&A*, 370, 447  
 Backman D. E., Paresce F., 1993, in Levy E. H., Lunine J. I., eds, *Protostars and Planets III*. Univ. of Arizona Press, Tucson, p. 1253  
 Barrado y Navascues D., 1998, *A&A*, 339, 831  
 Barrado y Navascues D., Stauffer J. R., Hartmann L., Suchitra S. C., 1997, *ApJ*, 475 313  
 Benz W., Asphaug E., 1999, *Icarus*, 142, 5  
 Bohren C. F., Huffman D. R., 1983, *Absorption and Scattering of Light by Small Particles*. Wiley, New York  
 Bouchet P., Manfroid J., Schmider F. X., 1991, *A&AS*, 91, 409  
 Britt D. T., Consolmagno G. J., 2001, *Icarus*, 152, 134  
 Brown M. E., Blake G. A., Kessler J. E., 2000, *ApJ*, 543, L163  
 Cameron A. G. W., 1997, *Icarus*, 126, 126  
 Campo Bagatin A., Petit J.-M., 2001, *Icarus*, 149, 210  
 Campo Bagatin A., Petit J.-M., Farinella P., 2001, *Icarus*, 149, 198  
 Canup R. M., Asphaug E., 2001, *Nat*, 412, 708  
 Cellino A., Michel P., Tanga P., Zappalà P., Paolicchi P., dell’Oro A., 1999, *Icarus*, 141, 79  
 Cooper J. F., Christian E. R., Johnson R. E., 1998, *Adv. Space Res.*, 21, 1611  
 Davis D. R., Chapman C. R., Weidenschilling S. J., Greenberg R., 1985, *Icarus*, 62, 30  
 Davis D. R., Weidenschilling S. J., Farinella P., Paolicchi P., Binzel R. P., 1989, in Binzel R. P., Gehrels T., Matthews M. S., eds, *Asteroids II*. Univ. of Arizona Press, Tucson, p. 805  
 Dermott S. F., Jayaraman S., Xu Y.L., Gustafson B. S., Liou J. C., 1994, *Nat*, 369, 719  
 Dohnanyi J. S., 1969, *J. Geophys. Res.*, 74, 2531  
 Dominik C., Tielens A. G. G. M., 1997, *ApJ*, 480, 647  
 Durda D. D., Dermott S. F., 1997, *Icarus*, 130, 140  
 Durda, D. D., Flynn G. J., 1999, *Icarus*, 142, 46  
 Durda, D. D., Greenberg R., Jedicke R., 1998, *Icarus*, 135, 431  
 Farley K. A., 1995, *Nat*, 376, 153  
 Flynn G. J., Moore L. B., Klöck W., 1999, *Icarus*, 142, 97  
 Fujiwara A., Tsukamoto A., 1980, *Icarus*, 44, 142  
 Fujiwara A., Cerroni P., Davis D. R., Ryan E., di Martino M., Holsapple K., Housen K., 1989, in Binzel R. P., Gehrels T., Matthews M. S., eds, *Asteroids II*. Univ. of Arizona Press, Tucson, p. 240  
 Gault D. E., Shoemaker E. M., Moore H. J., 1963, *NASA Tech. Note*, D-1767  
 GIBLIN I., Martelli G., Farinella P., Paolicchi P., Di Martino M., Smith P. N., 1998, *Icarus*, 134, 77  
 Gladman B., Kavelaars J. J., Nicholson P. D., Loredò T. J., Burns J. A., 1998, *AJ*, 116, 2042  
 Greaves J. S. et al., 1998, *ApJ*, 506, L133  
 Greenberg J. M., 1998, *A&A*, 330, 375  
 Grogan K., Dermott S. F., Durda D. D., 2001, *Icarus*, 152, 251  
 Hahn J. M., Malhotra R., 1999, *AJ*, 117, 3041  
 Hartmann W. K., 1988, *Lunar Planet. Sci.*, XIX, 451  
 Harvey P. M., Jefferys W. H., 2000, *ApJ*, 538, 783  
 Heiken G., 1975, *Rev. Geophys. Space Phys.*, 13, 567  
 Holland W. S. et al., 1998, *Nat*, 392, 788  
 Holland W. S. et al., 2002, *ApJ*, submitted  
 Holsapple K. A., 1993, *Ann. Rev. Earth Planet. Sci.*, 21, 333  
 Holsapple K. A., 1994, *P&SS*, 42, 1067  
 Housen K. R., Holsapple K. A., 1990, *Icarus*, 84, 226  
 Housen K. R., Holsapple K. A., 1999, *Icarus*, 142, 21  
 Housen K. R., Wilkening L. L., 1982, *Ann. Rev. Earth Planet. Sci.*, 10, 355  
 Jayawardhana R., Fisher S., Hartmann L., Telesco C., Piña, R., Giovanni F., 1998, *ApJ*, 503, L79  
 Jewitt D. C., 1999, *Ann. Rev. Earth Planet. Sci.*, 27, 287  
 Johnson R. E., Cooper J. F., Lanzerotti L. J., Strazzulla G., 1987, *A&A*, 187, 889  
 Jurac S., Johnson R. E., Donn B., 1998, *ApJ*, 503, 247  
 Kalas P., Larwood J., Smith B. A., Schultz A., 2000, *ApJ*, 530, L133  
 Kenyon S. J., 2002, *PASP*, 114, 265  
 Kenyon S. J., Bromley B. C., 2001, *AJ*, 121, 538  
 Kenyon S. J., Bromley B. C., 2002, *AJ*, 123, 1757  
 Koerner D. W., Sargent A. I., Ostroff N. A., 2001, *ApJ*, 560, L181  
 Kokubo E., Ida S., 1995, *Icarus*, 114, 247  
 Kokubo E., Ida S., 1996, *Icarus*, 123, 180  
 Kokubo E., Ida S., 2000, *Icarus*, 143, 15  
 Krivov A. V., Mann I., Krivova N. A., 2000, *A&A*, 362, 1127  
 Lachaume R., Dominik C., Lanz T., Habing H. J., 1999, *A&A*, 348, 897  
 Lagrange A. M., Backman D. E., Artymowicz P., 2001, in Mannings V., Boss

- A. P., Russell S. S., eds, *Protostars and Planets IV*. Univ. of Arizona Press, Tucson, p. 639
- Laor A., Draine B. T., 1993, *ApJ*, 402, 441
- Li A., Greenberg J. M., 1997, *A&A*, 323, 566
- Li A., Greenberg J. M., 1998, *A&A*, 331, 291
- Liseau R., 1999, *A&A*, 348, 133
- Lissauer J. J., 1993, *ARA&A*, 31, 129
- Lissauer J. J., Stewart G. R., 1993, in Levy E. H., Lunine J. I., Matthews M. S., eds, *Protostars and Planets III*. Univ. of Arizona, Tucson, p. 1061
- Love S. G., Ahrens T. J., 1996, *Icarus*, 124, 141
- Luu J. X., Jewitt D. C., 1998, *ApJ*, 502, L91
- Luu J. X., Jewitt D. C., Trujillo C., 2000, *ApJ*, 531, L151
- McKay D. S., Swindle T. D., Greenberg R., 1989, in Binzel R. P., Gehrels T., Matthews M. S., eds, *Asteroids II*. Univ. of Arizona Press, Tucson, p. 617
- Mennesson B., Mariotti J. M., 1997, *Icarus*, 128, 202
- Michel P., Benz W., Tanga P., Richardson D. C., 2001, *Sci*, 294, 1696
- Murray C. D., Dermott S. F., 1999, *Solar System Dynamics*. Cambridge University Press, Cambridge
- Nakamura A., Fujiwara A., 1991, *Icarus*, 92, 132
- Opik E. J., 1951, *Proc. R. Irish Acad.*, 54A, 165
- Ozernoy L. M., Gorkavyi N. N., Mather J. C., Taidakova T. A., 2000, *ApJ*, 537, L147
- Paolicchi P., Verlicchi A., Cellino A., 1996, *Icarus*, 121, 126
- Petit, J.-M., Farinella P., 1993, *Cel. Mech. Dyn. Astron.*, 57, 1
- Plets H., Vynckier C., 1999, *A&A*, 343, 496
- Pollack J. B., Hollenbach D., Beckwith S., Simonelli D. P., Damon P., Roush T., Fong W., 1994, *ApJ*, 421, 615
- Ryan E. V., Davis D. R., Giblin I., 1999, *Icarus*, 142, 56
- Song I., Caillaut J.-P., Barrado y Navascués D., Stauffer J. R., 2001, *ApJ*, 546, 352
- Stern S. A., 1996, *A&A*, 310, 999
- Tanaka H., Inaba S., Nakazawa K., 1996, *Icarus*, 123, 450
- Tanga P., Cellino A., Michel P., Zappalà V., Paolicchi P., dell'Oro A., 1999, *Icarus*, 141, 65
- Telesco C. M. et al., 2000, *ApJ*, 530, 329
- Toth I., 2000, *A&A*, 360, 375
- Trujillo C. A., Luu J. X., Bosh A. S., Elliot J. L., 2001, *AJ*, 122, 2740
- Veverka J., Thomas P., 1979, in Gehrels T., ed., *Asteroids*. Univ. of Arizona Press, Tucson, p. 628
- Weidenschilling S. J., 1977, *Ap&SS*, 51, 153
- Weidenschilling S. J., Spaute D., Davis D., Marzari F., Ohtsuki K., 1997, *Icarus*, 128, 429
- Wetherill G. W., Stewart G. R., 1993, *Icarus*, 106, 190
- Wiechert U., Halliday A. N., Lee D.-C., Snyder G. A., Taylor L. A., Rumble D., 2001, *Sci*, 294, 345
- Wilner D. J., Holman M. J., Kuchner M. J., Ho P. T. P., 2002, *ApJ*, 569, L115
- Wurm G., Blum J., 1998, *Icarus*, 132, 125
- Wyatt M. C., 1999, PhD thesis, Univ. Florida
- Wyatt M. C., 2001, in Dent W. R. F., ed., *Techniques For The Detection Of Planets And Life Beyond The Solar System*. Occasional Rep. R. Obs. Edinburgh, 17
- Wyatt M. C., Dermott S. F., Telesco C. M., Fisher R. S., Grogan K., Holmes E. K., Piña, R. K., 1999, *ApJ*, 527, 918
- Wyatt M. C., Dent W. R. F., Greaves J. S., Holland W. S., 2000, in Penny A. et al., eds, *Proc. IAU symp., Planetary Systems in the Universe: Observation, Formation and Evolution*. Astron. Soc. Pac., San Francisco, in press
- Zappala V., Cellino A., dell'Oro A., Migliorini F., Paolicchi P., 1996, *Icarus*, 124, 156
- Zuckerman B., 2001, *ARA&A*, 39, 549

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